

# Effects of land use on soil CO<sub>2</sub> flux in the Paramo de Guerrero, Colombia

Efectos del uso del suelo sobre el flujo de CO<sub>2</sub> del suelo en el Páramo de Guerrero, Colombia

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## ABSTRACT

The Andean paramo is an important global carbon sink and has a fundamental ecological function of capture, regulation and supply of water resources. The soil CO<sub>2</sub> efflux is a natural process through which the carbon is released into the atmosphere by molecular diffusion. The aim of this study was to establish the effect of different land use and soil managements practices over CO<sub>2</sub> efflux in the Paramo de Guerrero, using the soil respiration chamber technique. We evaluated five different land covers present in the Paramo de Guerrero (paramo vegetation, pasture, two tillage cover and potato crop). Our results show that soil respiration was lower in the paramo (0.42 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup>) than in the others land uses, probably due to the higher moisture content (57.1% on average). The tillage practices showed a primary physical effect, continued by the increase of the velocity of biological and chemical processes driven by soil microorganisms, such as microbial respiration and organic matter mineralization. This study demonstrates that moisture and soil temperature were not the main drivers of CO<sub>2</sub> flux in the conditions of the Paramo de Guerrero, but the agricultural management and the land use affect differentially the accumulation and release dynamics of soil organic carbon to the atmosphere.

**Key words:** agricultural soil management, carbon sinks, climate change, soil respiration, tillage.

## RESUMEN

El páramo es un importante sumidero de carbono global y además tiene funciones ecológicas fundamentales de captura, regulación y suministro del recurso hídrico. El flujo de CO<sub>2</sub> del suelo es un proceso natural a través del cual el carbono es liberado a la atmósfera por difusión molecular. El objetivo de este trabajo fue establecer los efectos de diferentes usos y prácticas de manejo del suelo sobre el flujo de CO<sub>2</sub> en las condiciones del Páramo de Guerrero, usando la técnica de cámara de respiración portátil de suelo. Se evaluaron cinco diferentes coberturas del suelo presentes (vegetación de páramo, pastura, dos lotes arados recientemente y cultivo de papa). Nuestros resultados muestran que la respiración en suelo fue más baja en el páramo (0.42 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup>) que en otros usos del suelo, probablemente debido al alto contenido de humedad (57,1% en promedio). Las prácticas de labranza mostraron un efecto físico principal, seguido por el aumento de la velocidad de procesos biológicos y químicos conducido por los microorganismos del suelo, tales como la respiración microbiana y la mineralización de la materia orgánica. Este estudio demuestra que la humedad y la temperatura del suelo no son los controladores principales del flujo de CO<sub>2</sub> del suelo en condiciones del páramo de Guerrero, pero el manejo agrícola y los cambios del uso del suelo afectan diferencialmente la dinámica de acumulación y liberación del carbono orgánico del suelo a la atmósfera.

**Palabras clave:** manejo agrícola del suelo, sumideros de carbono, cambio climático, respiración de suelos, arado.

## Introduction

The dense alpine vegetation called the “Paramo” have a fundamental ecological function of capture, regulation and water resources provision as an important socio economic service of water supply in the Andean ecosystem for different uses. However in the last decades, agricultural activities in these high mountain ecosystems have been dramatically intensified (Buytaert *et al.*, 2006a), becoming the main cause of paramos degradation (Poulenard y Podwojewski, 2000), altering and decreasing their potential for delivery of ecosystem services provision and regulation. Agriculture at the paramos cause a damage in soil

physical properties due to the breakdown and degradation of soil aggregates, which changes negatively the structure and the porosity, and generate a decreasing in the capacity of water retention and CO<sub>2</sub> capture, traits associated to high contents of organic matter, that play a key role in the function of water regulation (Buytaert *et al.*, 2006b) and carbon sequestration.

In general, the soil is an important carbon (C) sink and under the paramo conditions this storage capacity is high, due to the low redox potential, water content and humid climate. The carbon dioxide (CO<sub>2</sub>) flux is a natural process by which C is released into the atmosphere through soil

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respiration (Rs) (Fiener *et al.*, 2012; Oyonarte *et al.*, 2012; Pacific *et al.*, 2008; Riveros-Iregui *et al.*, 2008), which includes two components, autotrophic and heterotrophic respiration (Balthazar *et al.*, 2015; Riveros-Iregui *et al.*, 2008). Rs has a high spatial variability associated to changes in land use, geomorphological position and climate conditions (Davidson, 2006; Oyonarte *et al.*, 2012). Few studies have researched the influence of geomorphological position as a determining factor in soil respiration (Fiener *et al.*, 2012; Riveros-Iregui *et al.*, 2008). Likewise, the effects of land use changes over CO<sub>2</sub> flux have been rarely analyzed (Wang y Fang, 2009). Some authors report the influence of vegetation cover (Iqbal *et al.*, 2008; Xiao-Zeng *et al.*, 2013), and the effects of the land use change from natural vegetation to agriculture with different management practices altering the CO<sub>2</sub> efflux (Liu *et al.*, 2011; Motavalli *et al.*, 2000; Oyonarte *et al.*, 2012; Zhang *et al.*, 2013). In the surrounding paramos from Bogota D.C., there are no studies that associate the soil respiration response to the different land use changes and soil management practices.

The aim of this study was to establish the effect of different land use and soil managements practices, as also edaphoclimatics and geomorphologics conditions, on CO<sub>2</sub> efflux in the Paramo de Guerrero located at the municipality of Tausa, department of Cundinamarca, Colombia. Thirty six measurement sites of Rs were taken in a watershed of approximately 1,500 ha, to answer the following questions: 1) Are the edaphoclimatic conditions and geomorphological

position the main drivers of soil CO<sub>2</sub> flux? and 2) Are the land cover change and use the main regulators of soil CO<sub>2</sub> efflux and C in the Paramo de Guerrero?

## Materials and methods

### Location

This study was conducted in the watershed of Guandoque's river with a total extension of 1,508 ha. that fall into the Neusa reservoir located at 7 km to the southeast (Fig. 1). The watershed is located at 65 km to the north of Bogota in the municipality of Tausa, Cundinamarca department, Colombia. The climate is classified from a very cold to extremely cold, between semihumid, humid and very humid, with bioclimatic paramo's floor, low land paramo and andean highlands (IGAC, 1997).

We selected thirty six measuring sites in transects (Fig. 2) under five different vegetation covers based on cover and landuse classification map from the Guandoque's watershed (Peña, 2015). The cover presented in the study area included, clean pasture (*Pennisetum clandestinum*), potato crops (*Solanum tuberosum*), Paramo coverage (includes *Espeletia* sp., shrubland or dense grassland *Cavendishia nitida*, *Clusia multiflora*, *Drimys granadensis*. According to Morales *et al.* (2007) other species are presented in the area *Miconia* sp., *Weinmannia* sp., *Clusia* sp., *Ilex kunthiana* y *Drimys granadensis*. Two field plots with different soil preparation management were included in this study.

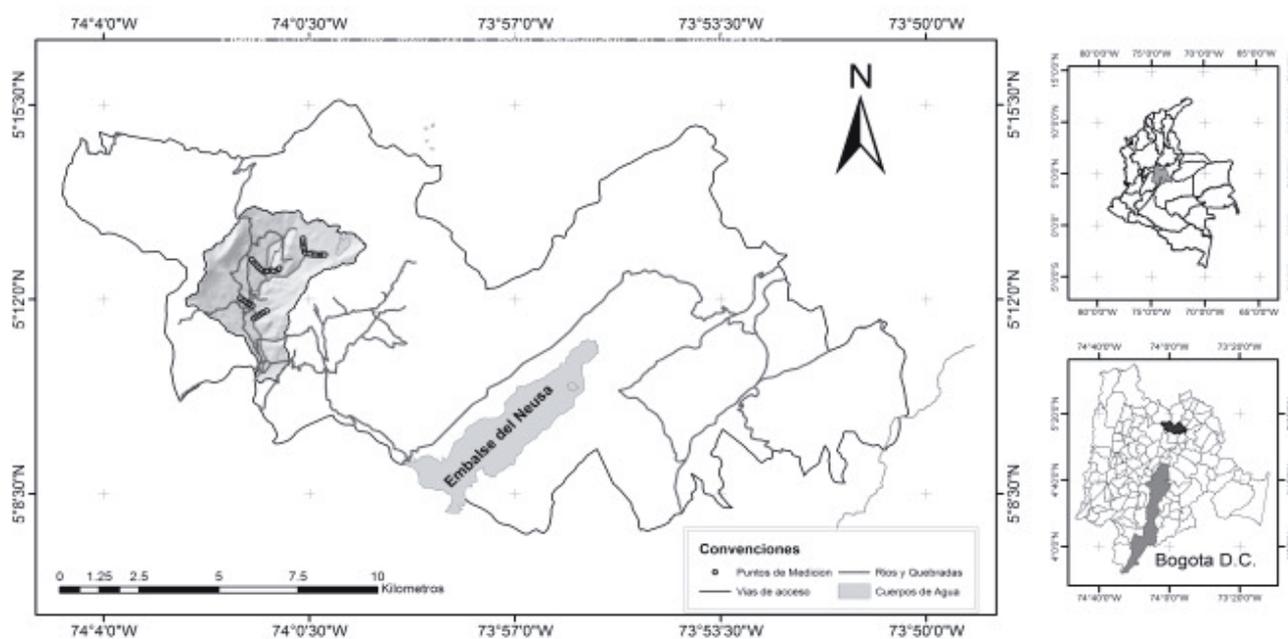
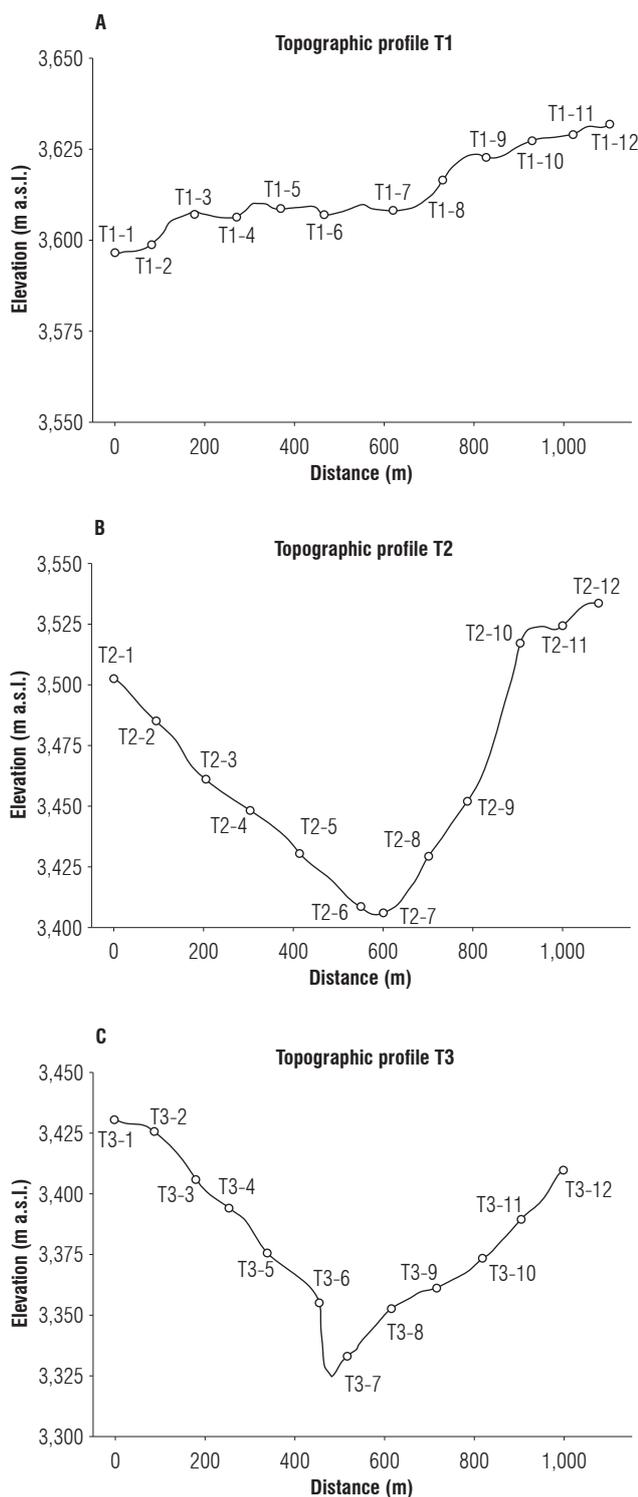


FIGURE 1. Location of the study area: municipality of Tausa, Cundinamarca (Colombia).



**FIGURE 2.** Topographic profiles of the measuring transect. A. Transect 1; B. Transect 2, and C. Transect 3.

**Transect 1 (T1):** paramo coverage, without agricultural use.

**Transect 2 (T2):** tillage field plot (with previous fallow coverage, soil tillage was done with 1 pass of disc plow to a depth

of 30-40 cm when starting the experiment, followed by 1 rotavator pass to a depth of 10-20 cm approximately one month after), pasture and paramo coverage.

**Transect 3 (T3):** tillage field plot (with previous potato crop coverage, 1 pass of rotavator to a depth of 10-20 cm, about three months after the potato crop harvesting, and 1 month after the beginning), potato crop and pasture.

**Reference plot:** fragmented forest coverage without agricultural intervention, situated in the same watershed and in the same altitudinal range of T2 and T3 (the reference plot was used just to compare the total organic carbon contents).

### Edaphoclimatic and geomorphological characteristics

Soil sampling and description was done by means of detailed observations on excavated holes of 50 cm (width) × 50 cm (high) × 50 cm (deep). The internal characteristics were described (Peña, 2015), and as a result we found majority of soil are derived from volcanic ashes, order Andisol. For the external characteristics of soil, the methodology proposed by the Instituto Geográfico Agustín Codazzi was followed (IGAC, 2014). The transect 1 (T1) was located in the highest elevation area of the watershed, in artesa relief type, with a difference in elevation of 30 m, approximately, among the different sampling sites, while for the transects T2 and T3 the differences among sampling sites were 140 and 100 m, respectively (Fig. 2), both transects in terraces relief type and fluvio-lacustrine hills. Although there are differences in elevation, T2 and T3 are located in the same climate, classified as very cold (temperature between 8 and 12°C), while transect T1, is located in an extremely cold weather (temperature between 4 and 8°C).

### Field measurements

The experiment was conducted from January 12, 2014 until March 27, 2014. The  $R_s$  was measured in the 36 sampling sites with three replicates per plot, every measurement was taken every 2-7 d, for a total of 1,620  $CO_2$  efflux data.  $R_s$  measurements were carried out with a portable infrared gas analyzer EGM-4 with a 0-2.000 ppm range (Environmental Gas Monitor, PP Systems, UK) and the soil respiration chamber SRC-1 with a calibration range of 0-9,99  $g\ CO_2\ m^{-2}\ h^{-1}$  (PP Systems, UK). Measurements were performed after 120 s of  $CO_2$  concentration stabilization in the SRC-1 chamber with flux adjustment done by the quadratic regression (Peña, 2015; Riveros-Iregui *et al.*, 2008). On each monitored site the soil moisture ( $m^3\ m^{-3} \times 100$ ) was recorded with the Hydrosense II (Cambell Scientific) portable measurement system and the soil temperature was registered at 12 cm depth (REOTEMP). The  $R_s$  was measured in the

36 sampling sites of 1×1 m approximately, marked with a colored ribbon to avoid disturbance.

### Total organic carbon

In addition to the description, disturbed samples were collected to measure the total organic carbon (TOC) at different depths, 5, 20 and 50 cm, by using the dry total combustion method (elemental analyzer), the TOC was also quantifying for the forest fragmented cover use as reference.

### Data Analysis

To contrast the null hypothesis, which states that the independent variable transects (geomorphological position) has no effect on the response variables (CO<sub>2</sub> flux, soil temperature and soil moisture), the multivariate analysis of variance MANOVA was used as an extension of the analysis of variance (ANOVA) for some response variables (Peña, 2002). The Bonferroni test was used to calculate

multiple comparison of means from different levels of transect factor.

## Results

### Soil CO<sub>2</sub> efflux variability associated with edaphoclimatic conditions

In addition to the differences in altitude, the slopes recorded in T1 are lower than those of T2 and T3 (Fig. 3), these slope conditions acting together with climate generated the differences in soil moisture. The soil moisture content (Hs) and soil temperature (Ts) showed a strong negative linear correlation (-0.78), and the transect T1 was the one with the highest moisture and lowest temperature (Fig. 3).

Soil CO<sub>2</sub> efflux in T1 was lower (0.42 g CO<sub>2</sub> m<sup>-2</sup>h<sup>-1</sup>) than the flux of T2 and T3 (0.48 y 0.43 g CO<sub>2</sub> m<sup>-2</sup>h<sup>-1</sup>, respectively), also had the lowest variation coefficient (T1=30%, T2=46% and T3=43%), related with the homogeneity in cover type in T1, while T2 and T3 had 3 different covers (Fig. 4). Transect T1 registered the highest moisture content (57%) and the lowest average temperature (11.2 °C) between the three measuring transects. While the moisture recorded in transects T2 and T3 was 27% and 31% respectively and the averages temperature in both transects was 14°C.

The effect of the independent variable “transect” on the dependent variables CO<sub>2</sub> efflux, Ts and Hs was significant, and based on the results of Wilks’ Lambda and Pillai’s Trace (Pr>F = <0.0001) test, the null hypothesis was rejected, which indicated the influence that geomorphological position has on edaphoclimatic conditions (Ts and Hs) for each site and CO<sub>2</sub> efflux.

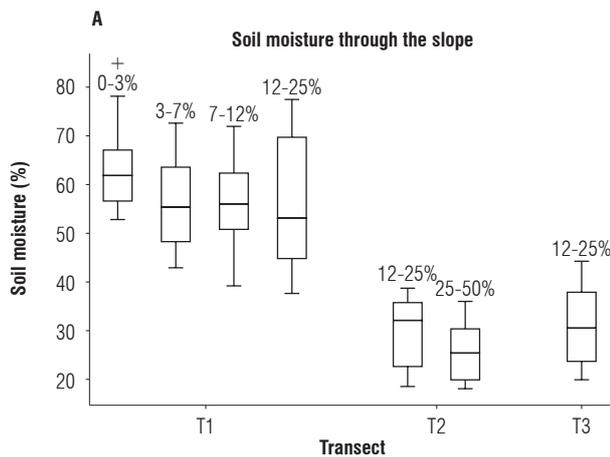


FIGURE 3. Soil moisture content variability, in relation to the slope of transects T1, T2 and T3 in the Paramo de Guerrero (Colombia).

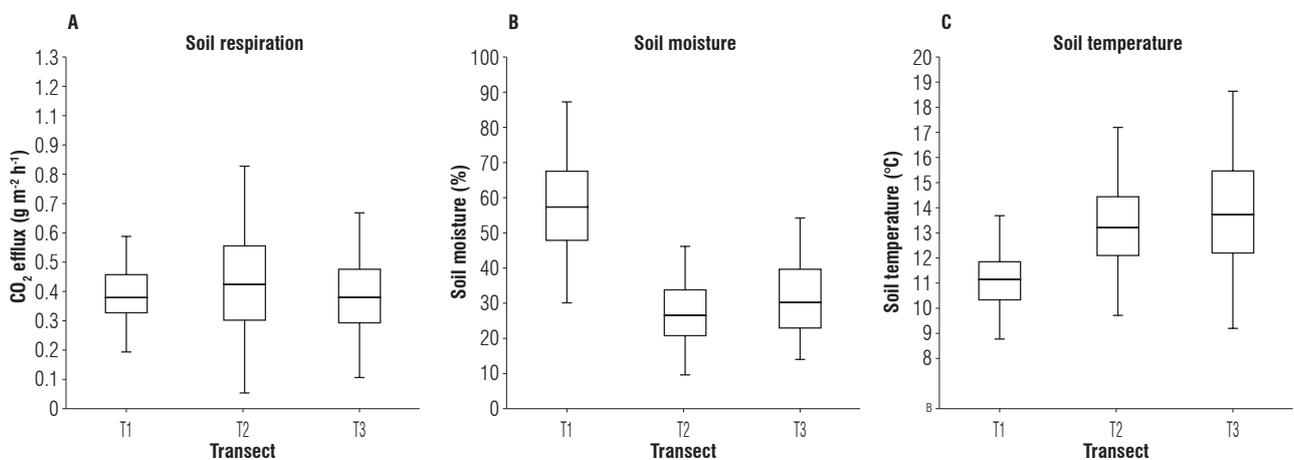


FIGURE 4. CO<sub>2</sub> flux variability, soil moisture content and temperature in T1, T2 and T3 transects in the Páramo de Guerrero, Colombia.

With the Bonferroni test the effects among transects were contrasted for each soil variable. Table 1 shows that CO<sub>2</sub> flux in T1 transect had significant differences with the T2 transect, but not with the T3 transect, while T2 showed significant differences with T3 (Fig. 4). In Table 2 it was observed that there were significant difference regarding moisture content among transects, while Tab. 3, shows that there were only significant differences in soil temperature between T1 and the other two transects (Fig. 4), nevertheless, there was no difference in Ts between T2 and T3 transects.

**TABLE 1.** Bonferroni test for Contrast among transects for CO<sub>2</sub> flux variable.

Transects comparison	Differences between means	Simultaneous confidence limits 95%		
T2 - T3	0.05036	0.00284	0.09787	*
T2 - T1	0.05127	0.00195	0.10059	*
T1 - T3	-0.00091	-0.05036	0.04853	

\*  $P \leq 0.05$

**TABLE 2.** Bonferroni test for contrast among transects for Hs variable.

Transects comparison	Differences between means	Simultaneous confidence limits 95%		
T1 - T3	26.565	23.729	29.401	*
T1 - T2	29.936	27.107	32.764	*
T3 - T2	3.37	0.645	6.096	*

\*  $P \leq 0.05$

**TABLE 3.** Bonferroni test for contrast among transects for Ts variable.

Transects comparison	Differences between means	Simultaneous confidence limits 95%		
T2 - T3	-0.4216	-0.9233	0.0802	
T2 - T1	2.342	1.8212	2.8628	*
T1 - T3	-2.7636	-3.2857	-2.2415	*

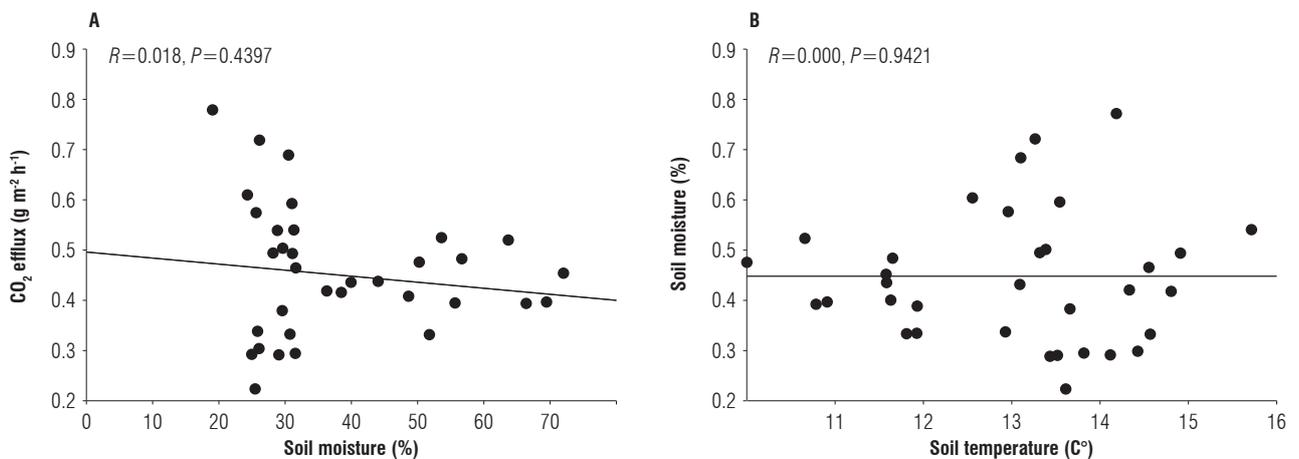
\*  $P \leq 0.05$

Although there were significant differences regarding the edaphoclimatic conditions among transects, the moisture ( $P=0.4397$ ) and temperature ( $P=0.9421$ ) variables did not show any linear relation with the CO<sub>2</sub> flux variable (Fig. 5).

### Effect of land use and cover

There were not found significant differences in CO<sub>2</sub> flux among covers, although the emissions ranges might have showed a CO<sub>2</sub> flux variability between them (Fig. 6). Fallow coverage in T2 was the one that registered a greatest CO<sub>2</sub> flux ( $0.59 \text{ g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ), while the Paramo cover in T2 recorded the lowest CO<sub>2</sub> flux ( $0.31 \text{ g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ). The pasture cover in T2 and T3, did not have significant differences in CO<sub>2</sub> flux. The tillage field plots in T2 and T3 ( $0.34 \text{ g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ) reported contrasting results of CO<sub>2</sub> flux despite the similarity in external conditions of the plots (bared soils), but had differences.

As the tow tillage plots has two different background agromomic management systems, was possible to identify two different periods of the CO<sub>2</sub> flux behavior, after each soil tillage practices, with both tillage instrument (Fig. 7). The A (raising) period corresponded to the significant increase in the CO<sub>2</sub> flux, after some tillage practice was applied, and period E (stabilization) shows a CO<sub>2</sub> flux reduction and regulation. In T2 disk plowing (fallow) it was identifying two A periods of significant CO<sub>2</sub> flux raise, and two E periods following the A periods of CO<sub>2</sub> flux increase (Figure 7a), that coincide with the two tillage pass done on this plot. On the other hand, in the disk plow plot apply after potato in T3 there was only one considerable A period CO<sub>2</sub> flux, associated to the tillage event, followed by an E period at the end of the study (Fig. 7b).



**FIGURE 5.** Non significant linear regression of CO<sub>2</sub> flux with soil moisture in the three transects. A. no significant linear regression between CO<sub>2</sub> flux and B. Temperature in the three transects.

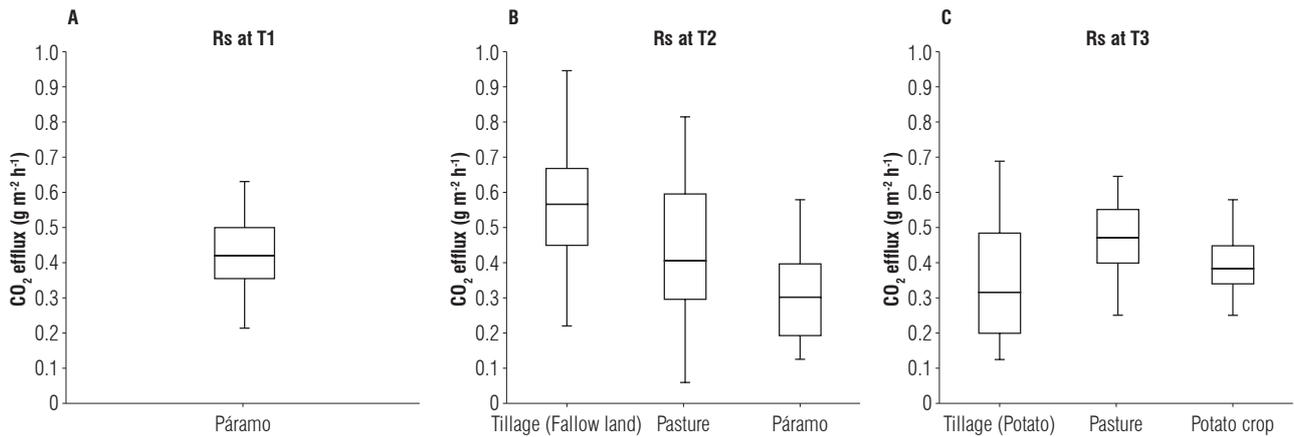


FIGURE 6. Soil CO<sub>2</sub> flux variability for each cover. A. Transect T1; B. Transect T2 and C. Transect T3 at the Guerrero's Páramo.

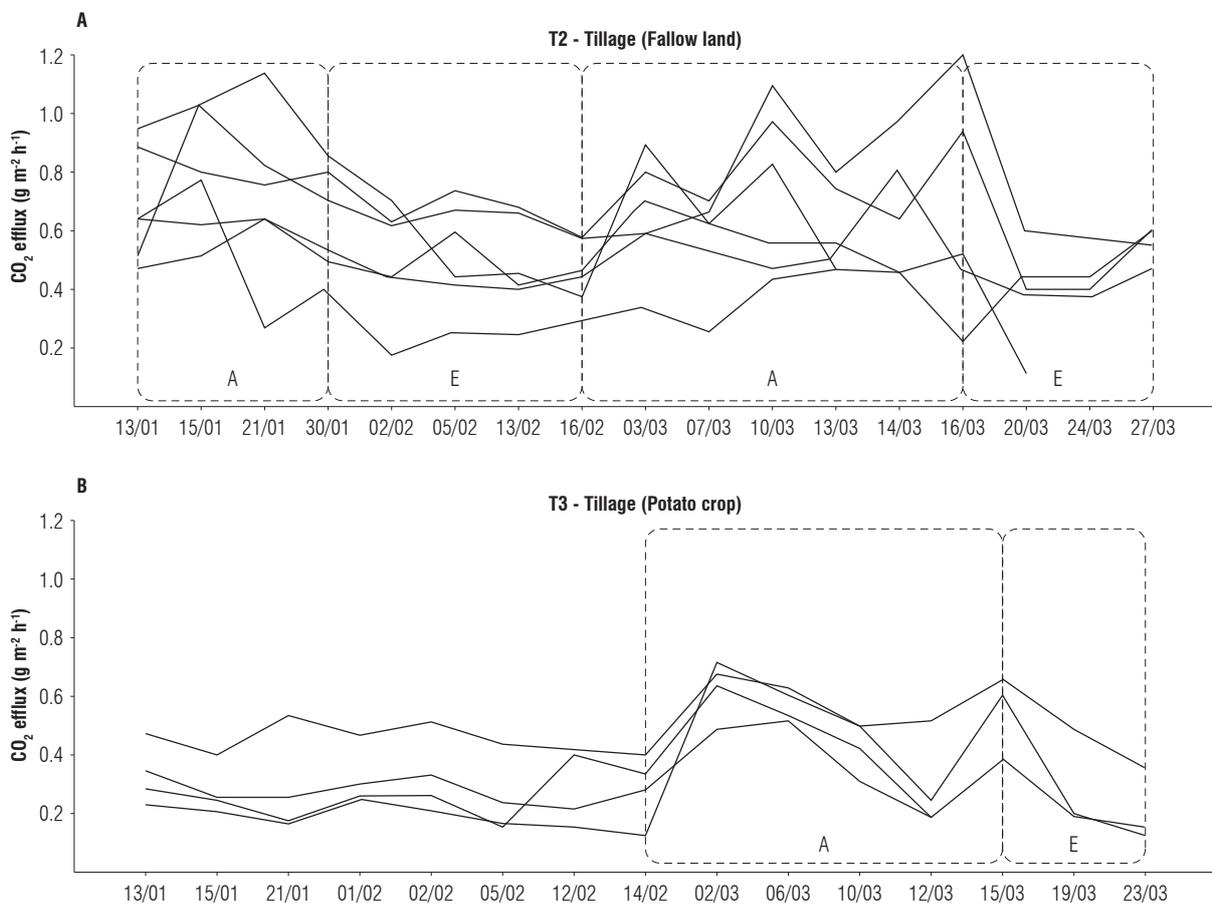


FIGURE 7. Effects of soil management on CO<sub>2</sub> flux in T2 transects, two tillage labors with a disk plowing after a fallow cover, and the effects of one tillage labor on T3, in the Paramo de Guerrero. Each labor represents a raising period (A) of CO<sub>2</sub> efflux. Every line represents the CO<sub>2</sub> efflux through all the experiment for each plot with tillage cover. A. T1 and B. T2.

During the first A period there was a maximum CO<sub>2</sub> total flux between 2 and 4 kg ha<sup>-1</sup> h<sup>-1</sup> in T2 after the first disk plow coming from a long period of rest from agricultural activity. Otherwise, after the second tillage, the maximum increase of CO<sub>2</sub> flux was lower (1-3 kg ha<sup>-1</sup> h<sup>-1</sup> of CO<sub>2</sub>). In

T3, after the first tillage, the maximum CO<sub>2</sub> flux was lower (2 to 3 kg ha<sup>-1</sup> h<sup>-1</sup> CO<sub>2</sub>) compared to T2.

The intensive tillage in annual and semi-annual crops accelerate the process of physical, chemical and biological

soil degradation, reducing the carbon stock due to the soil loss caused by the different types of erosion and oxidation (mineralization), especially in tropical and subtropical areas (Silva *et al.*, 1994; Souza *et al.*, 2014). The different types of cover had an effect on the total organic carbon content (Fig. 8), the reference cover, which not have any kind of severe anthropogenic disturbance, was used to compare the effect of land use changes on total organic carbon content.

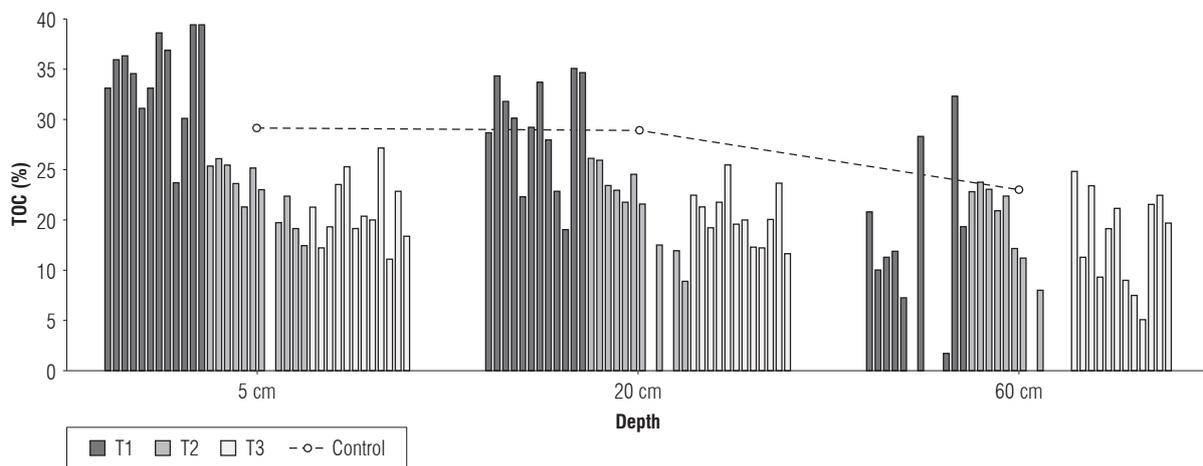
In figure 8, TOC contents of T1, T2 and T3 transects were observed for each of the measuring sites, the missing bars refer to sites in which the sampling could not be performed (due to plot conditions, plot close to crop (T2-8) or the soil from the sampling site did not have the requirement depth). It was observed that the TOC contents in T1 (26.25% average) were higher than those recorded in T2 (18.28% average) and T3 (17.37% average), and even higher than those obtained in the reference plot, which could be related to the

different covers and management practices, and also with the edaphoclimatic conditions of each site, in areas that presented higher moisture content at 5 and 20 cm depth the organic matter accumulation processes are preferred over the mineralization processes (Fig. 9).

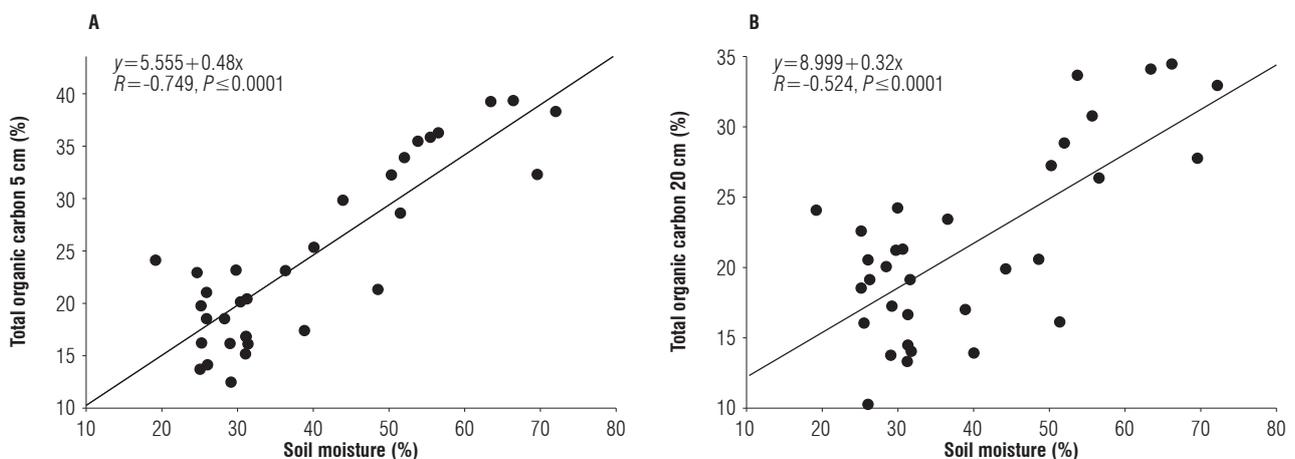
The cover change effects and management in T1 and T2 transect caused a TOC estimated loss of 5% for the first 5 cm depth, however to 20 cm depth TOC losses are highest, and reach an mean of 8% (Fig. 8) compared with the reference plot. In 50 cm depth, the TOC contents were similar in T2, T3 and the reference.

## Discussion

Although different authors report that the soil moisture content is the most important controlling CO<sub>2</sub> flux factor (Pacific *et al.*, 2008; Riveros-Iregui *et al.*, 2008;



**FIGURE 8.** Effects of the type of soil management on total organic carbon content in the T1, T2 and T3 transect, with forest reference in the Paramo de Guerrero (Control).



**FIGURE 9.** Relationship between the Hs and the TOC to the 5 cm (A) and 20 cm depth (B).

Riveros-Iregui *et al.*, 2008) and others indicated that temperature is the main driver of this process (Campos, 2006; Iqbal *et al.*, 2008; Subke y Bahn, 2010). In the present study no linear or polynomial relation between CO<sub>2</sub> flux with Hs or Ts was found. But, it was observed that in wet conditions and low soil temperatures there was a positive relation (Fig. 9) between the moisture content and the TOC accumulation. The results endorsed the common knowledge that the Paramo is a carbon storage ecosystem, which needs to be protected and preserve in order to help mitigate the effects of climate change.

The investigation showed no significant difference in the CO<sub>2</sub> flux among the transects. The results of CO<sub>2</sub> flux are similar to those reported by Shahidi *et al.* (2014) in two agricultural soils (Alfisol and Mollisol) with ranges of 0.4 to 1.2 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> under conventional tillage systems and non till, the greatest flows were obtained in the Mollisol order probably due to the highest content of organic carbon, compared to the Alfisol order. Liu *et al.* (2011) also reported similar values in an Oxisol order of a bamboo plantations in the subtropical region of China.

Transects T1 and T2 have the same paramo cover, without any human intervention. The difference found between the CO<sub>2</sub> flux values (Fig. 6), are explained based on the differences in geomorphological traits and slope in T2. In the paramo cover of transect T1, the slopes are lower than 25% and soils are moderately deep, greater than 50 cm, compared to the paramo cover existing in the T2 transect, where slopes between 25% and 50% were found, with surface soils with less than 50 cm thickness. These conditions affect the process of organic materials accumulation (main input in CO<sub>2</sub> flux).

The immediate increases in CO<sub>2</sub> flux after conventional tillage practices are the result of CO<sub>2</sub> physical release (Fig. 7) stored in the pores of the soil as a result of microbial activity, these increases in the release of CO<sub>2</sub> are the evidence of the land use change and management practices (Álvaro-Fuentes *et al.*, 2007). The increased accumulation of CO<sub>2</sub> fluxes were greater than the reported by Ellert y Janzen (1999) who obtained increases of 3.3 kg ha<sup>-1</sup> h<sup>-1</sup> of CO<sub>2</sub>, during the first 6 hours after tillage and increases of 0 to 0.5 kg ha<sup>-1</sup> h<sup>-1</sup> of CO<sub>2</sub> during the first week of monitored, related to aeration and residual mixture that increase microbial activity in the rhizosphere and soil aggregates. During these stages on which microbial activity is favored, the increase of CO<sub>2</sub> flux is mediated by the Rh on the Ra (Shahidi *et al.*, 2014).

The effects of tillage practices were also reflected in the TOC content in T2 and T3 transects (Fig. 8) with an history of agricultural management, compared to T1 and the reference plot without agricultural intervention. The reduction in TOC content is related mainly to the annual and semester crops, such as potatoes (*Solanum tuberosum*) and some vegetables, or rotation such as pasture or fallow (Wang *et al.*, 2006). The different types of land use and management produced losses in the TOC contents. This results are corroborated by Lal (2003) who found that the management practices that commonly control the organic C storage are tillage, chemical fertilization, crop rotation and crop density. The reductions in the TOC content were found at depths of 5 and 20 cm, associated with the direct influence of the different tillage practices, since the optimum depth for crop development, especially potatoes is 20 to 30 cm.

The reduction in TOC content is in the range of 5% to 5 cm deep and 8% at 20 cm depth (Fig. 8), compared to the reference fragmented forest cover. The reductions in the TOC are lower than those reported by Motavalli *et al.* (2000) who found losses between 20 and 40% of the organic carbon in soils in which forest cover was change to agricultural crops. Similar results were reported by Dube *et al.* (2009) with losses of 44% of the organic carbon content in 10 cm soil surface, when natural forest was change to pine plantations. Constant tillage techniques, particularly the use of moldboard plow in association with other types of plow (chisel, disk, rigid chisel), may have greatest negative impact on several soil quality indicators, in particular those associated with TOC (Karlen *et al.*, 2013). After tillage, the mineralization of the organic matter increased rapidly due to changes in the amount of water, air and temperature effects in soil. A reduction in TOC concentrations minimizes soil resilience, decreases its quality, reduces its productivity and buffering capacity (Lal y Bruce, 1999). When tillage practices are applied to fresh plant material with some degree of decomposition is incorporated into the soil, during oxygenation phase generated by the rupture and exposure of aggregates in the surface horizons, the process of organic matter mineralization is favored, allowing higher release of greenhouse gases (GEI) (CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O) (Álvaro-Fuentes *et al.*, 2007; Datta *et al.*, 2013; Kutsch *et al.*, 2010). The optimization of fertilizer management, irrigation and tillage operations or the adoption of non-till practices could contribute to reducing GHG emissions and the recovery of 60 to 70% of the organic matter content that has been lost from soil (Chen *et al.*, 2015; Datta *et al.*, 2013).

## Conclusions

This investigation showed that moisture and soil temperature were not the main drivers of CO<sub>2</sub> efflux in the climatic and edaphoclimatic conditions of the Paramo de Guerrero. Nevertheless, the slope and the geomorphological position did contribute in the soil moisture content and temperature. Although moisture content was positively related with the TOC storage, this not affected the C dynamics emission into the atmosphere as CO<sub>2</sub>, on the contrary high soil moisture means more C accumulation on soil, and high soil moisture is one of the principal characteristics of Andean paramos, converting them in big sink of C in soil.

The effects of land use and cover change on soil CO<sub>2</sub> flux into the atmosphere indicated that conventional tillage practices, the tillage intensity and the management background, physically affected the CO<sub>2</sub> releases previously stored by microbiological processes in the soil pores, besides of increasing the microorganism's activity on newly added organic materials or on the organic material stored in the soil.

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