

CHARACTERIZATION OF NATURAL MICROCOSMS OF ESTUARINE MAGNETOTACTIC BACTERIA

CARACTERIZACIÓN DE MICROCOSMOS NATURALES DE BACTERIAS MAGNETOTÁCTICAS ESTUARINAS

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ABSTRACT: To date, no complete study of magnetotactic bacteria's (MTB) natural microcosms in estuarine or tropical environments has been reported. Besides, almost all the studies around magnetotactic bacteria have been based on fresh waters away from the Equator. In this work, we focused the experimental region at the Equator and present a comprehensive mineralogical and physicochemical characterization of two estuarine bacterial microcosms. The results show that mineral lixiviation in the sediments may be an important factor in the solubilization of elements required by magnetotactic bacteria. Specifically, we show that clinocllore, phlogopite, nontronite, and halloysite could be among the main minerals that lixiviate iron to the estuarine microcosms. We conclude that nitrate concentration in the water should not be as low as those that have been reported for other authors to achieve optimal bacteria growth. It is confirmed that magnetotactic bacteria do not need large amounts of dissolved iron to grow or to synthesize magnetosomes.

KEY WORDS: Magnetotactic bacteria (MTB), magnetosome, microcosm, estuary

RESUMEN: No se ha reportado ningún estudio completo sobre microcosmos naturales de bacterias magnetotácticas (MTB) en estuarios o ambientes tropicales. Además, casi todos los estudios sobre las bacterias magnetotácticas se han desarrollado en aguas dulces alejadas del ecuador. Este trabajo se desarrolla sobre el ecuador y reporta una caracterización mineralógica y fisicoquímica detallada de dos microcosmos bacterianos estuarinos. Los resultados muestran que la lixiviación de minerales en los sedimentos puede ser un factor importante en la solubilización de elementos requeridos por las bacterias magnetotácticas. Específicamente, que el clinocloro, flogopita, nontronita y haloisita pueden estar entre los minerales más importantes en la lixiviación de hierro a los microcosmos estuarinos. Se concluye que la concentración de nitrato en el agua no debe ser tan baja como se ha reportado para lograr un crecimiento bacteriano óptimo. Las bacterias magnetotácticas no necesitan grandes cantidades de hierro disuelto para su crecimiento ni para la síntesis de magnetosomas.

PALABRAS CLAVE: Bacterias magnetotácticas (MTB), magnetosomas, microcosmos, estuario.

1. INTRODUCTION

Magnetotactic bacteria are microorganisms of the bacteria domain, whose directional swimming behavior is affected by the Earth's geomagnetic and external magnetic fields [1-2]. This property is known as *magnetotaxis* [3-4]) and occurs mainly due to the presence of magnetic nanocrystals (generally of magnetite [Fe₃O₄], or greigite [Fe₃S₄]) that shape an intracellular, single-magnetic-domain and membrane-bounded structure known as *magnetosome* [2, 4]. This property is generally assumed to facilitate the bacteria in its finding and maintaining a favorable position in

vertical chemical gradients in stratified environments [5, 6]. Currently, a wide morphology variety of MTB has been reported, such as coccus, bacillus, vibrio, spirillum, and multicellular aggregates [2, 7, 8].

Many authors have studied the natural environment of different species of MTB, searching for a strategy to obtain large amounts of magnetic nanocrystals [4, 9, 10]. The objective of those studies was to identify the most important physico-chemical factors that are involved in the growth of MTB populations and the synthesis of magnetosomes [1, 4, 10]. Some of these chemically dissolved factors, are the iron (total Fe, Fe²⁺, and Fe³⁺), sulfates, nitrates in the solution,

and dissolved oxygen (DO) [1, 4, 11, 12]. Furthermore, all the reported MTB are anaerobic or strict-microaerophilic [7, 12], and mesophilic [1, 10]. Another factor that could be important in those processes is the microbial ecology of the natural environments of MTB, but it has not yet been studied thoroughly [1].

In this research, different spectroscopic techniques and chemical analyses were used to characterize two estuarine MTB microcosms, situated in the tropical waters of the Caribbean Sea. We report the variations in soil mineralogy, water composition, and some physico-chemical parameters of the MTB-environments. This information may serve as a guide to elucidate potential mineral donors that contribute to the formation of magnetosomes and other intracellular bodies.

2. MATERIALS AND METHODS

2.1 Sampling zone

The samples were taken in two different estuaries. Cispatá Bay (BC) (9°21' - 9°25' North latitude, 75°45' - 75°50' West Longitude) and the Caimanera Bog (CC) (9°25' north latitude, 75°41' east longitude) (Fig. 1). Both estuarine systems are located in the Morrosquillo Gulf in the Colombian Caribbean Sea and are conformed by a great variety of mangrove swamps that shelter an abundant population of marine and estuary species.

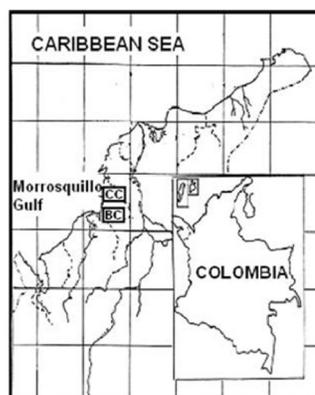


Figure 1. Location of sampling zone: Cispatá Bay (BC) and Caimanera Bog (CC) in the Morrosquillo Gulf, Colombia.

The water-sediment samples were taken in the Oxidic-Anoxic Transition Zone (OATZ), in northern regions of both estuaries. A HANNA oxymeter was used to locate the OATZ in the water column. The depth of sampling in BC and CC was around 2 m and 1.5 m, respectively.

2.2 Spectroscopy

The spectroscopy techniques used to study the composition of the sediments were: X-ray diffraction (XRD), Fourier transform infrared spectroscopy (FTIR), and Mössbauer spectroscopy (MS). Before the analyses, the sediments were dried at room temperature and milled to pass through a 200 Tyler mesh.

2.2.1 X-ray Diffraction

The XRD analyses were carried out in Panalytical X'pert Pro MDP equipment, in a 2θ range of 10° to 70° , at a speed of $0.02^\circ/s$, using Cu $K\alpha$ radiation with a current of 40 mA and 45 kV. The results were analyzed with the diffraction software DIFFRAC^{plus} 2000 and the data base PDF2.MDI.

Relative abundance of the minerals in the sediments was qualitatively estimated based on the height of the peaks in the spectra.

2.2.2 Fourier transform infrared spectroscopy

The FTIR analyses were carried out in a Spectrum One Perkin Elmer Spectrophotometer, operating by transmittance between 4000 and 450 cm^{-1} . The pellets were made with milled sediments and KBr.

2.2.3 Mössbauer spectroscopy

Wissel Mössbauer equipment was used in transmission and constant acceleration mode for Mössbauer spectra (MS) acquisition. The equipment was operated with ^{57}Co in a rhodium matrix.

2.3 Water analysis and physico-chemical parameters

2.3.1 Water analysis

The codes of the methods used for the water analyses refer to the STANDARD METHODS FOR EXAMINATION OF WATER AND WASTEWATER [13]. Total alkalinity (mg/L CaCO_3) (2320B); alkalinity to phenolphthalein (mg/L CaCO_3) (2320B); chlorides (mg/L Cl^-) (4500-CI B); phosphates ($\text{mg/L PO}_4^{3-}\text{-P}$) (4500-P D); total phosphorus (mg/L P) (4500-P D); nitrates (mg/L N-NO_3^-) (4500- NO3- A); nitrites (mg/L N-NO_2^-) (4500- NO2-B); ammoniacal nitrogen ($\text{mg/L NH}_3\text{-N}$) (4500- NH3 C.D); total nitrogen

(mg/L N) (4500 - NORG); organic nitrogen (mg/L N) (4500 - NORG); total, ferrous and ferric iron (3500 Fe); sulfates (4500 SO_4^{2-}).

2.3.2 Physico-chemical parameters

We used a digital thermometer, HACH HQ40d pH-meter, SCHOOTT Eh-meter and HI 9143 HANNA oxymeter, to measure temperature, pH, Eh, and dissolved oxygen (DO), respectively. All the measurements were *in situ*. Average salinity was calculated in the laboratory by the evaporation of a fixed volume of estuary water. Salinity was calculated as an average.

2.4 MTB and magnetite nanocrystals presence

MTB from BC and CC were magnetically isolated using the glass recipient described by Ulysses Lins *et al.* (2003) [14]. Magnetotaxis were confirmed by optical microscopy (Olympus CX31). Magnetite nanocrystals were detected by electronic microscopy (Phillips, Tecnai G²), operating at 200kV, and identified by Energy-dispersive X-ray spectroscopy (EDX). Additionally, MTB population was counted in natural BC and CC samples using a BOECO Neubauer chamber.

3. RESULTS

3.1 X-ray Diffraction

Figure 2 shows the XRD spectra for BC and CC sediments. Quartz appears as the most abundant mineral in both sediments. No magnetite or gregite were detected by XRD in BC or in CC. However, MTB (and its magnetite magnetosomes) presence in the systems was confirmed by optical and electronic microscopy.

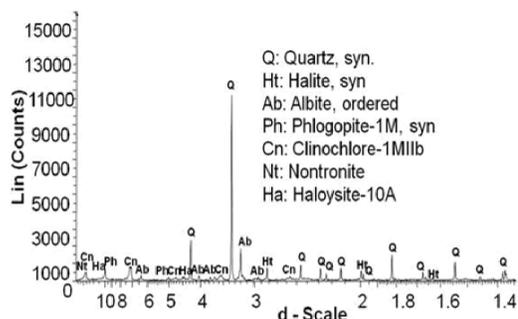
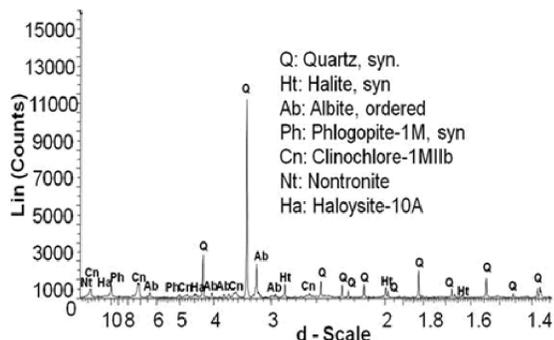


Figure 2. XRD spectrum of a sediment sample from BC (up) and CC (down). The graphics are on d-scale.

3.2 Fourier transform infrared spectroscopy

Figure 3 shows the FTIR spectra of BC and CC.

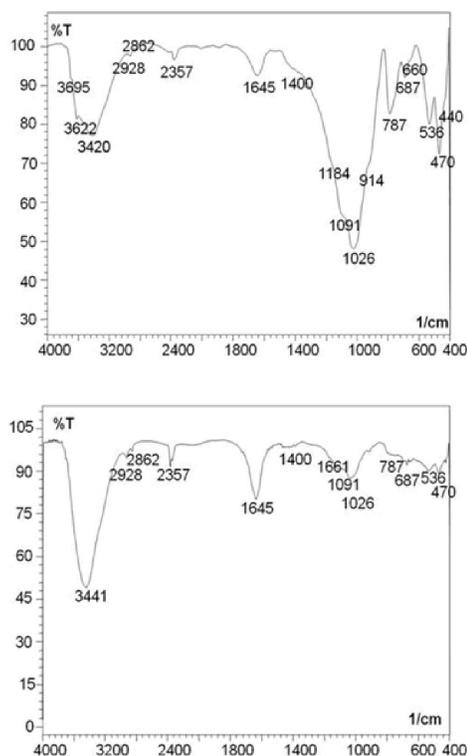


Figure 3. FTIR spectrum of a sediment sample from BC (up) and CC (down).

Bands in 3420 cm^{-1} and 1645 cm^{-1} (Fig. 3, left), and in 3441 cm^{-1} and 1645 cm^{-1} (Fig. 3, right), are commonly associated with vibrations of the H-O bonds in water molecules [15, 16]. Bands in 2928 cm^{-1} and 2862 cm^{-1} are reported as vibrations of C-H bonds in organic compounds [16]. The little band in 2357 cm^{-1} is associated with vibrations in the CO_2 molecule [15, 17]. Bands on the right

hand side of both spectrums are commonly associated to silicates. Specifically, bands in 1091 cm⁻¹, 787 cm⁻¹ and 687 cm⁻¹, are reported as vibrations of Si-O bonds in quartz structure [15, 18] and in 3695 cm⁻¹, 3622 cm⁻¹, and 914 cm⁻¹ such as Si-O vibrations in the structure of halloysite [16]. Bands in 1400 cm⁻¹, 1184 cm⁻¹, 1026 cm⁻¹, 660 cm⁻¹, 536 cm⁻¹, 470 cm⁻¹, and 440 cm⁻¹ are reported as vibrations of Si-O bonds in many silicates [15, 19]. Based on DRX analyses, those bands which are in the spectrum of BC sediments (Fig. 3, left) can be associated with quartz, albite, phlogopite, clinocllore, nontronite, and halloysite; and in CC sediments (Fig. 3, right) to quartz, albite, phlogopite, clinocllore, nontronite, and gypsum [15, 20]. It is possible that a lower band, near to 580 cm⁻¹, indicates the presence of a small amount of magnetite [21, 22] in both estuaries. In this case, those bands would be overshadowed by the highest bands.

3.3. Mössbauer spectroscopy

Figures 4 and 5 show the Mössbauer spectra, of sediment samples from BC and CC, respectively. Additionally, Tables 1 and 2 show the quadrupole splitting (Qs), isomeric shift (Is) and contribution percentage of the minerals in the sediments [23].

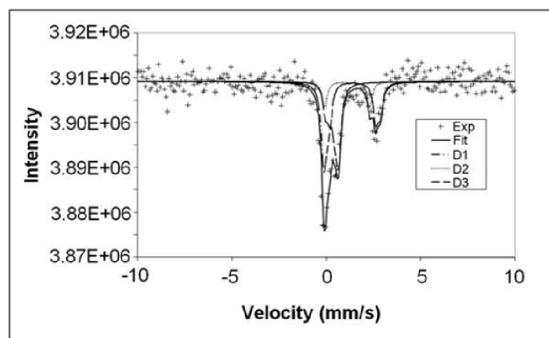


Figure 4. Mössbauer spectrum of a sediment sample from BC. It shows the experimental data (Exp), three doublets (D1, D2 and D3) and the fit.

| | Qs (mm/s) | Is (mm/s) | Iron phase | % Phase | Contributing minerals |
|----|-----------|-----------|------------------|---------|---|
| D1 | 0.8 | 0.35 | Fe ⁺³ | 53 | Clinocllore Nontronite Halloysite |
| D2 | 2.4 | 1.2 | Fe ⁺² | 13 | Clinocllore Nontronite |
| D3 | 2.4 | 1.5 | Fe ⁺² | 34 | Clinocllore Nontronite Phlogopite |

Table 1. Mineral Mössbauer parameters that contribute to the iron phases in BC sediments.

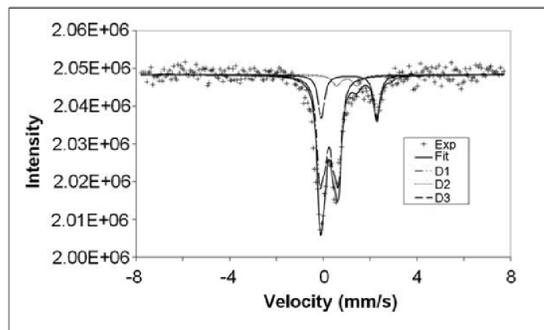


Figure 5. Mössbauer spectrum of a sediment sample from CC. It shows the experimental data (Exp), three doublets (D1, D2 and D3) and the fit.

| | Qs (mm/s) | Is (mm/s) | Iron phase | % Phase | Contributing minerals |
|----|-----------|-----------|------------------|---------|---|
| D1 | 0.8 | 0.35 | Fe ⁺³ | 69 | Clinocllore Nontronite Halloysite |
| D2 | 0.9 | 1.1 | Fe ⁺² | 8 | Clinocllore Nontronite |
| D3 | 2.4 | 1.2 | Fe ⁺² | 23 | Clinocllore Nontronite Phlogopite |

Table 2. Mineral Mössbauer parameters that contribute to the iron phases in CC sediments.

It was not possible to determine the relative percentage of each phase, since the hyperfine parameters of the minerals present are very similar and a superposition of the subspectra appears, making its discrimination impossible. This overlap is reflected by the presence of very wide doublets in the spectra. Because of this problem, an analysis based on distributions of hyperfine parameters, using the DISTRI program was chosen [24]. Therefore, the values presented are average values for the isomeric shift and quadrupole splitting.

Table 3 lists the minerals detected by XRD and MS spectroscopy in BC and CC estuaries and their relative abundance in the sediments. Quartz is the most abundant mineral in both estuaries. Halite and albite are also present in both estuaries, but they are much less abundant than quartz. The other minerals are present in the sediments as traces.

| Mineral name | CC MS | CC XRD | BC MS | BC XRD |
|-----------------|-------|--------|-------|--------|
| Quartz | nd | +++++ | nd | +++++ |
| Halite | nd | + | nd | + |
| Albite | nd | + | nd | + |
| Phlogopite | d | t | d | t |
| Clinocllore | d | t | d | t |
| Nontronite | d | t | d | t |
| Pyrite | d | t | nd | nd |
| Halloysite | nd | nd | d | t |
| Gypsum | nd | t | nd | nd |
| Ferroactinolite | nd | t | nd | nd |

Table 3. Relative abundance of minerals in BC and CC estuary sediments. Very high abundance (+++++), high (++++), moderately high (+++), moderate (++), low (+), trace (t), detected (d), not detected (nd).

3.4 Water analysis and physico-chemical parameters

3.4.1 Water analysis

Table 4 shows the results of water analyses for BC and CC samples, respectively.

Table 4. Water analysis of BC (up) and CC (down).

| Parameters | Results |
|---|---------|
| Total alkalinity (mg/L CaCO ₃) | 236 |
| Alkalinity to phenolphth. (mg/L CaCO ₃) | 0 |
| Chlorides (mg/L Cl ⁻) | 21193 |
| Phosphates (mg/L PO ₄ ³⁻ -P) | 0.511 |
| Total Phosphorus (mg/L P) | 0.938 |
| Nitrates (mg/L N-NO ₃ ⁻) | 2.93 |
| Nitrites (mg/L N-NO ₂ ⁻) | 0.04 |
| Ammoniacal nitrogen (mg/L NH ₃ -N) | 5.04 |
| Total nitrogen (mg/L N) | 5.46 |
| Organic nitrogen (mg/L N) | 4.20 |
| Total iron (mg/L) | 0.958 |
| Ferrous iron Fe ²⁺ (mg/L) | 0.285 |
| Ferric iron Fe ³⁺ (mg/L) | 0.673 |
| Sulphates (mg/L) | 1300 |
| Parameters | Results |
| Total alkalinity (mg/L CaCO ₃) | 180 |
| Alkalinity to phenolphth. (mg/L CaCO ₃) | 0 |
| Chlorides (mg/L Cl ⁻) | 21243 |
| Phosphates (mg/L PO ₄ ³⁻ -P) | 0.212 |
| Total Phosphorus (mg/L P) | 0.345 |
| Nitrates (mg/L N-NO ₃ ⁻) | 2.55 |
| Nitrites (mg/L N-NO ₂ ⁻) | 0.08 |
| Ammoniacal nitrogen (mg/L NH ₃ -N) | 3.08 |
| Total nitrogen (mg/L N) | 6.02 |
| Organic nitrogen (mg/L N) | 5.04 |
| Total iron (mg/L) | 3.308 |
| Ferrous iron Fe ²⁺ (mg/L) | 1.371 |
| Ferric iron Fe ³⁺ (mg/L) | 1.937 |
| Sulphates (mg/L) | 840 |

Both sites are very similar in this feature. Nevertheless, there are some differences to point out. The BC microcosm has a higher alkalinity than that of CC. The BC waters have more than twice the amount of phosphates and a slightly higher amount of ammoniacal nitrogen than CC. And finally, total iron (in Fe²⁺ and Fe³⁺ phases) are significantly higher in the BC microcosm than in CC.

3.4.2 Physico-chemical parameters

Temperature, pH, Eh, DO, and the average salinity from BC and CC estuaries are shown in Table 5.

| | T (°C) | pH | Eh (mV) | DO (mg/L) | Average salinity |
|----|--------|------|---------|-----------|------------------|
| BC | 29 | 7.51 | 130 | 2.10 | 4.07% ± 0.37% |
| CC | 30 | 7.49 | 118 | 2.61 | 4.64% ± 0.34% |

Table 5. Physico-chemical parameters in BC and CC estuaries.

Both estuaries have a neutral pH and a low redox potential, with very low amounts of dissolved oxygen (microaerophilic MTB). Those conditions are common for MTB [6, 9]. As well as all the MTB reported in the literature, the bacteria of this study are mesophilic.

3.5 MTB and magnetite nanocrystals presence

Figure 6 shows two characteristic cocci MTB from BC and CC estuaries. Each bacterium has four magnetosome chains/cell. Each magnetosome have around seven magnetite nanocrystals.



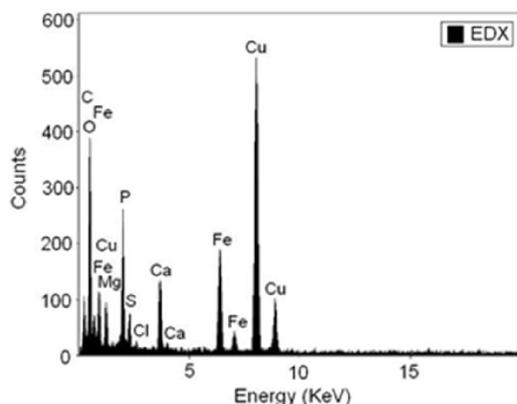


Figure 6. (Top) Cocci MTB abundant in BC and CC estuaries. Four magnetosomes can be seen in each cell. (Bottom) EDX of a magnetite nanocrystal in a magnetosome of MTB shown in the left image (star).

Figure 6 also presents the EDX spectrum taken from the star-point showed in the left image. The iron and oxygen percentages in the EDX spectrum are very similar to those reported for magnetite (72.36% Fe; 24.64% O) [15]. The Cu detected is from the mesh.

On the sampling day, BC microcosms sheltered a MTB population of $(7.1 \pm 0.1) \times 10^5$ MTB/ml; and CC microcosms sheltered a MTB population of $(7.6 \pm 0.1) \times 10^5$ MTB/ml. The number of non-magnetic microorganisms was much greater than the MTB.

4. DISCUSSION

Quartz and halite, the most abundant minerals in BC and CC sediments, are reported as the most common minerals in estuarine and coastal sediments [24, 25]. Although it is usually considered that the concentration of quartz does not have a significant effect on the population dynamics of estuarine microorganisms, the halite concentration does play a major role in estuarine's microbial dynamics. Low-salinity (<10%) estuarine areas have been identified as regions of severe physiological stress for both freshwater and marine organisms [26].

Clinochlore, nontronite, and halloysite were detected in BC sediments by X-ray diffraction. Furthermore, FTIR bands at 3695, 3622 and 914 cm^{-1} , seem to confirm the presence of halloysite [15, 16]. FTIR bands at 1400, 1184, 1026, 660, 536, 470, and 440 cm^{-1} (Si-O vibrations) support the idea that those clay minerals

are present in BC and CC sediments. This is also supported by the XRD and Mössbauer analysis. The Mössbauer spectrum of BC sediments is interpreted as the contribution of those silicates to form two doublets of ferrous iron (47%) and one of ferric iron (53%). Surely these minerals play an important role in the microcosm's iron cycle, mainly as a source of lixiviated ferric iron. That is why the mineralogical composition of the microcosm's soils is thought to be an important issue in the study of the natural microcosms of MTB.

The sediments of the CC estuary can be analyzed in the same way. In this case, clinochlore, nontronite, and phlogopite are the contributing minerals in the Mössbauer spectrum. Pyrite was also detected by XRD, but it is unlikely that this is lixiviated under the reported physicochemical conditions for CC. The reductive environment and neutral pH induce the pyrite nucleation rather than its lixiviation. As in the BC case, all minerals were detected by XRD and supported by FTIR spectroscopy. There were also three doublets in the Mössbauer spectrum. The average values of isomeric shift and quadrupole splitting suggest that Fe^{3+} doublet in CC sediments is due to nontronite and clinochlore contributions (69%); and Fe^{2+} doublets are due to clinochlore, nontronite (23%), and phlogopite (8%). Nevertheless, phlogopite is much less abundant in the sediments than clinochlore and nontronite. Ferroactinolite is also in CC sediments, and could be a source of lixiviated iron, but the Mössbauer analysis indicates that this is much less abundant than the other contributing minerals. We conclude that the iron lixiviated from clinochlore and nontronite, appears as the main source of the iron used by the MTB from CC microcosms, for cell growth and magnetosomes synthesis.

Several previous reports have implied that MTB are involved in the reduction of iron and it has been suggested that they could derive energy from this process [2, 7]. Flies *et al.* (2005) [9] found MTB in a zone where abundant solid phase Fe^{3+} was present, and thus it was considered that it might serve as an electron acceptor. Nevertheless, authors like Kim *et al.* (2005) [10], did not find overwhelming evidence of whether or not an increase on the iron concentration is necessary for MTB growth. In the present research, we have found that the CC estuary has a greater iron concentration than BC estuary, but there appears to be

no significant difference between the MTB populations in these microcosms.

Biogenic magnetite nanocrystals were detected in both MTB microcosms. Nevertheless, magnetite was not detected in any of the estuary sediments. It is possible that a weak and overshadowed signal of magnetite appears in both FTIR spectra, but it was not detected by XRD spectroscopy, and the characteristic sextet of magnetite [23] in Mössbauer spectra did not appear either. However, it must be considered that these spectroscopy techniques require an amount of magnetite that is unlikely to be found in natural sediments in its biogenic form [27]. Also, magnetite is found in sediments in the form of nanocrystals of superparamagnetic nature, making it difficult to be detected either by MS or XRD. More appropriate techniques such as magnetic measurements should be used. Cummings *et al.* (2000) [27] noted that even if the population density of MTB was relatively high, the magnetite contribution by magnetotactic bacteria could not by itself explain the dominant magnetic character of the sediments.

The EDX spectrum indicated not only the presence of iron and oxygen (for the magnetite), but also phosphorus (P), sulphur (S), calcium (Ca), chlorine (Cl), magnesium (Mg), and carbon (C). Certainly those elements appear in the spectrum due to the cell membrane and the cytosol interference. The dark bodies inside the MTB in Fig. 6 (left) have been reported as intracellular vesicles of sulphur, phosphate, and/or chlorine [4, 28]. Those intracellular bodies can also contribute to the EDX spectrum shown above. It is considered that a generic cell has an average composition of: 50% C, 20% O, 14% N, 8% H, 3% P, 1% S, 1% K, 1% Na, 0.5% Ca, 0.5% Cl, 0.5% Mg, 0.2% Fe, and 0.3% of other trace elements [29]. This explains the presence of these elements in the EDX spectrum in Fig. 6 (right). Also, the fact that Fe, P, Ca, and Mg, appear in larger amounts in MTB than in a generic cell, is further evidence of the existence of the intracellular MTB-vesicles. It is important to point out that all these elements are available in BC and CC microcosms.

Oxygen is one of the elements that is thought to be restrictive to MTB population growth [7, 9, 12], especially to microaerophilic MTB. However, MTB have also proven to be very sensitive to other chemical gradients, such as sulfide [9, 10], dissolved iron [2, 7],

nitrites, and nitrates [9]. MTB, likes many mesophilic organisms, seems to be very tolerant to changes in temperature. According to this study, the MTB of the BC and CC estuaries are capable of being in a reductive environment (low Eh values) of neutral pH and a relatively high concentration of DO. Flies *et al.* (2005) [9] reported the absence of nitrates and nitrites as optimal to MTB growth. But, Tables 4 and 5 show values of nitrates near to 3mg/L in both estuaries, and yet the MTB population of BC and CC microcosms (around $7.4 \cdot 10^5$ MTB/ml) reaches similar values to those they reported (around $1.5 \cdot 10^6$ MTB/ml) [9]. On the other hand, the average concentration of dissolved iron in BC and CC waters is similar to that reported as optimal by Kim *et al.* (2005) [10] (near to 0 mg/L) and close to the minimal values reported by Simmons *et al.* (2004) [1] (between 0.558 mg/L – 12.566 mg/L). Although several researchers have published that magnetotactic bacteria needs a very stable and specific chemical gradient to growth [5], the findings of the present study show that MTB can adapt to variable environments like the constantly changing estuaries.

5. CONCLUSIONS

In spite of the intrinsic variability of estuary systems BC and CC, as ecosystems subjected to marine and continental influences, it is possible to characterize them by implementing the spectroscopy techniques and physico-chemical analyses used in this work. In fact, these results may serve as a guide for elucidating the possible origin of some intracellular elements in MTB, like Fe and S.

The natural population of MTB cannot by itself contribute to the sediments the necessary amount of biogenic magnetite to be detectable by XRD, FTIR, and Mössbauer spectroscopy. Accumulation techniques of MTB, like that proposed by Lins *et al.* (2003) [14], must be carried out to achieve that objective. The presence of magnetite in the sediments cannot be ruled out, and other more sensitive methods should be used.

It is very likely that the structural iron of biogenic magnetite is derived from the lixiviation of the minerals in the sediments. Clinocllore, nontronite, and halloysite could be among the main minerals that contribute iron to the magnetosomes synthesis in MTB. Phlogopite and ferroactinolite could also be among the

contributive minerals, but they are much less abundant in BC and CC estuaries than those mentioned above. It is unlikely that pyrite lixiviation occurs under the physicochemical conditions reported in the BC and CC microcosms. Although quartz appears as the most abundant mineral in the sediments, it does not seem to influence the MTB population dynamics.

If clinocllore, nontronite, halloysite, phlogopite, and ferroactinolite are actually among the main iron donors for magnetite biosynthesis, it would be necessary to assess the participation of MTB in the iron cycle of these aquatic environments. Indeed, some researchers are already taking advantage of the bioaccumulation capacities of MTB for the remotion of different contaminant compounds in wastewaters [30].

The development and maintenance of specific chemical gradients could be one of the major factors in the MTB population growth. Apparently, nitrate concentrations should not be as low as those that have been reported [9] in order to achieve an optimal MTB growth. The iron concentration in waters of BC and CC support the idea that MTB does not need large amounts of dissolved iron to survive or to synthesize magnetosomes. Further studies are needed to completely understand the relationships between MTB populations and their natural microcosms.

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