# Selection of earthquake ground motion accelerograms for the continental margin of Southeastern Brazil 

Cristian Yair Soriano Camelo ${ }^{a}$, Samuel Felipe Mollepaza Tarazona ${ }^{a}$, Maria Cascão Ferreira de Almeida ${ }^{b}$, Márcio de Souza Soares de Almeida ${ }^{a}$ \& Ricardo Garske Borges ${ }^{c}$<br>${ }^{a}$ Graduate School of Engineering - COPPE/UFRJ, Rio de Janeiro, Brazil.cysorianoc@coc.ufrj.br, samuelfelipe@coc.ufrj.br, almeida@coc.ufrj.br<br>${ }^{b}$ Polytechnic School of Engineering, - Poli/UFRJ, Rio de Janeiro, Brazil.mariacascao@poli.ufrj.br<br>${ }^{c}$ Petrobras Research and Development Center - CENPES, Rio de Janeiro, Brazil. ricardogarskeborges@gmail.com

Received: January $26^{\text {th }}, 2021$. Received in revised form: April 20 th, 2021 . Accepted: April $28^{\text {th }}, 2021$.


#### Abstract

Brazil is in an intraplate area of low to moderate seismicity, this means that few or no records of strong ground motions are available. Part of the site response analysis and seismic design of structures require the use of acceleration time-histories compatible with a specified target response spectrum. This study aims to utilize methodologies based on the use of existing earthquake records from a well-known database and synthetic accelerograms to obtain ground motions representative of the Brazilian Southeast Region, particularly in the offshore Campos Basin. Information from a probabilistic seismic hazard assessment performed in the interest area was employed as input to the methodologies applied in terms of target response spectrum and the dominant earthquake scenarios. Besides, the acceleration time-histories of two relatively recent earthquakes that occurred in the Brazilian Southeast were used to apply one of the approaches to obtain a synthetic spectrum compatible accelerogram.


Keywords: artificial accelerograms; earthquake accelerogram; ground motion; spectral matching.

## Selección de acelerogramas sísmicos para la margen continental de la región Sudeste de Brasil

## Resumen

Brasil se encuentra en una región intra-placa de baja a moderada sismicidad, esto quiere decir que hay una limitada disponibilidad de registros de acelerogramas representativos de la región. Parte de un estudio de respuesta local o del diseño sísmico de estructuras requiere del uso de registros aceleración-tiempo que sean compatibles con un espectro de diseño. El presente trabajo tiene como objetivo la aplicación de metodologías para obtener acelerogramas representativos de la región sudeste brasilera, en particular en la región offshore de la cuenca de Campos. Se utilizó la información del análisis probabilístico del riesgo sísmico desarrollado para la región de interés en este estudio para obtener el espectro de respuesta de diseño y para definir los escenarios sísmicos que generan mayor influencia. Adicionalmente, se utilizó la información de acelerogramas registrados a partir de dos sismos relativamente recientes ocurridos en la región sudeste brasilera, para la aplicación de una de las metodologías de obtención de acelerogramas sintéticos.

Palabras clave: acelerogramas artificiales; acelerogramas; sismos; ajuste espectral.

## 1. Introduction

Brazil is in a large mid-plate region considered one of the least seismically active continental areas in the world. Earthquakes with maximum magnitudes 5 and above occur with a return period of 4 years [1]. This means that recordings of
strong ground motions in the region are scarce. As alternative to obtain earthquakes to be applied in the seismic design of structures or site response analyses in such intraplate regions, different methodologies can be applied [32].

This paper aims to obtain a series of accelerograms representative of the earthquake scenarios in a region located

[^0] of Southeastern Brazil.. DYNA, 88(217), pp. 228-236, April - June, 2021.
in Southeastern Brazil, specifically in the Campos Basin Continental Slope, offshore of the State of Rio de Janeiro. The dominant earthquake scenarios and mean uniform hazard spectra used in this work for the region of interest were developed by [5]. Three methodologies were employed to obtain spectrum-compatible accelerograms. The first approach consisted of the use of real accelerograms obtained from the Pacific Earthquake Engineering Research (PEERNGA) database [58] and the application of spectral matching procedures. A second methodology of developing synthetic accelerograms generated from seismological source models $[45,46]$. And a third alternative, also based on synthetic accelerograms, using envelope shapes adjusted from real records of accelerograms in Brazil $[3,57]$.

### 1.1 Seismic hazard in Southeastern Brazil

Based on the distribution of seismicity, the main active seismic areas in Brazil were delineated by [1] as follows: (1) Southern part of the Guyana shield and middle Amazon basin; (2) a North-Southtrending zone along the Eastern border of the Amazon craton; (3) Northern part of the Borborena Province, in Northeastern Brazil around the Potiguar marginal basin; (4) the Porto dos Gauchos seismic zone; (5) a NE-SW zone in the Tocantins province possibly continuing towards the Pantanal Basin; (6) Southern Minas Gerais zone, in and around the southern tip of the São Francisco craton; and (7) the Southeast offshore zone with activity concentrated along the continental slope from the Pelotas Basin in the south to the Campos basin from $33^{\circ} \mathrm{S}$ to $20^{\circ} \mathrm{S}[2,3]$. The Campos Basin represents an important pole in Brazilian economy and is one of the most prolific Brazilian oil-producing basins [4]. Given the importance of the Campos Basin, moderate seismicity in the oceanic portion of the Southeast region (magnitude 5 every 1030 years and potential occurrences of magnitude up to 7) represent a hazard to the design of offshore infrastructure developments. The area of interest for the analyses presented in this work correspond to the continental margin of Southeastern Brazil. Specifically, the Campos Basin Continental Slope [6], offshore the state of Rio de Janeiro (Fig. 1).


Figure 1. Location of the site of interest (dark square). Source: The authors.

Two main sources of stress are associated in the seismicity pattern in the Brazilian Atlantic coast. A regional component of compressive stresses and a local component due to density contrast between continental/oceanic crusts and lithospheric flexure due to sediment load in the continental shelf [2]. The overall structural style in Campos Basin is detached, with a detachment surface at the base of the Aptian salt [7]. Below the detachment, the main structural features are horsts and grabens limited by steep normal faults [8,9]. The stress regime is variable with depth. Extensional regimes at depths of less than $1,500 \mathrm{~m}$, transtensional regimes at depths ranging from 1,500 to $3,500 \mathrm{~m}$, and strike-slip, for depths greater than $3,500 \mathrm{~m}$ [10].

### 1.2 Earthquake catalog

Brazilian earthquake data from historical and relatively recent instrumental records have been compiled through the earthquake catalog provided by the Seismology Center of the Institute of Astronomy Geophysics and Atmospheric Sciences of the University of São Paulo (IAG/USP). Further details on the earthquake catalog are presented [11]. Fig. 2 presents the location of the epicentres and magnitudes of the historical earthquakes covering a period from November 1720 to December 2019. It can be observed from Fig. 2 a concentration of epicenters in the Campos Basin, roughly across the continental slope [5].

## 2. Selection of earthquake ground motion accelerograms

For seismic hazard assessment, the definition of the ground motions represents a significant role. The most complete way to represent the ground motion is through seismograms that describe the variation with time of the accelerations, velocity, and displacements. It is usually displayed the variation of the accelerations with time (accelerogram), the velocities and displacements are obtained by integration.


Figure 2. Historical and instrumental seismic activity from 1720 to 2019.
Source: [5]

One common challenge of many intraplate regions is the limited amount of significant earthquake records due to the infrequent occurrence of these events. When two of the largest Brazilian earthquakes occurred in $1955\left(\mathrm{~m}_{\mathrm{b}} 6.2\right.$ on January 31 and $m_{b} 6.1$ on March 1) no Brazilian stations were operating [11]. Early instrumental data in the Brazilian earthquake catalog were obtained from international agencies (USGS and ISC) and using stations from other countries. Only in the late 1960s with the installation and in the 1970s instrumental records begin in earnest [12]. Further details related to the history of the seismological stations in Brazil can be found in [13]. The most recent boost in the seismological research in Brazil started in 2011 with the operation of the Brazilian Seismographic Network [11]. Given the scarcity of recorded accelerograms, an alternative to obtaining strong motion data consist of scaling and matching ground motions from other regions with local seismic hazard response spectra [14] or the generation of spectrum compatible accelerograms [15-18].

The accelerogram generation methods can be divided into the following categories: (1) deterministic methods, implementing a superposition of harmonic waves to match a target response spectrum [19]; (2) wavelet-based techniques [20-23]; (3) stochastic methods that consider the accelerogram as a Gaussian process [24-26]. In this work, three methodologies will be employed to obtain spectrum compatible ground motions. For the analyses, the Mean Uniform Hazard Response Spectra for the site of interest at the continental slope of the Campos Basin for a return period of 975 years was used as target spectrum (Fig. 3).

### 2.1 Method 1: seismic records from database

The first approach consisted of using recorded accelerograms from the Pacific Earthquake Engineering Research (PEER) Strong Motion Database. The PEER database consists of two databases: the NGA-East Ground Motion Database and the NGA-West Database [58]. The first constitutes the largest database of processed recorded ground motions in Stable Continental Regions (SRCs).


Figure 3. Uniform Hazard Response spectra for the site of interest at the continental slope of the Campos Basin for various return periods. Source: Adapted from [5].

The second includes a very large set of ground motions recorded worldwide of shallow crustal earthquakes in active tectonic regimes. As Brazil is located within a stable continental intraplate region, it was considered at a first instance to use the NGA-East database. However, from the results of a preliminary search, no records satisfied the criteria discussed in the following section. Then, the NGA-West database was used. A similar approach was used for the selection of earthquake ground accelerograms in Hong Kong, also an intraplate area of low to moderate seismicity [14].

### 2.1.1 Search and selection criteria

The PEER web application allows the user to load a target spectrum and a series of criteria to select the earthquakes (magnitude, fault type, significant duration, distance to rupture plane, average shear wave velocity of top 30 meters' site, and pulse characteristics). The ground motions can also be adjusted to reduce the Mean Squared Error (MSE) and increase the match between the target and recorded (scaled) spectra. A maximum limit for the scale factor of 2.0 was defined for the records that satisfied the search criteria. This value for the scale factor of 2.0 was proposed by [27].

From a disaggregation analysis, it is possible to obtain the contribution of different earthquake scenarios at the site of interest. The parameters considered are the magnitude intervals and distance ranges [28-30]. The moment magnitude and the source-to-site distance are two variables commonly used in the selection of accelerograms. The most relevant geophysical parameter for the selection of records is earthquake magnitude, as it strongly influences frequency content and duration of the motion [31]. Regarding the source-to-site distance, these effects have not been clearly established. Boomer and Acevedo [32] proposed that, in performing the selection of real records, the search window based on the magnitude should be as narrow as possible, and in terms of the distance, the range can be extended. [31] recommended the selection of records from events within 0.25 units of the target magnitude. [33] proposed select the time series, using records within 0.5 magnitude units of the design earthquake. For the current analyses, a search window in the earthquake's magnitude of $\pm 0.5 M_{w}$ was adopted, a similar range was used by [14]. The primary reason was that using a narrower search window yielded less or no results when consulting the records from the PEERNGA East database.

For the area of interest at the offshore Campos Basin, the most significant contribution comes from magnitudes 4.5 to 5.1 $M_{w}$ at distances between 40 and 100 km . Regarding large earthquakes (magnitudes up to $5.7 M_{w}$ ) it was defined a contribution at distances up to 200 km [5]. There are few events accurately located in the Brazilian earthquake catalog, and the source-to-site distance estimated was at less than about 15 km . In addition, [2] reported an average distance of 10 km for oceanic source zones. Subsequently, the minimum distance used for the search criteria was reduced from 40 to 10 km . Given the stress regime in the Campos Basin, the failure mechanism considered reverse and normal faults, however, to include larger amplitude motions reverse and reverse/oblique faults were accounted in the search criteria. The target site rock condition of the uppermost 30 m of the earth crust $\left(v_{s, 30}\right)$ was defined for $v_{s, 30}$ values higher than $800 \mathrm{~m} / \mathrm{s}$. With those criteria, three earthquake scenarios were defined, and the search parameters are shown in Table 1.

Table 1.
Parameters for the selection of accelerograms

| Earthquake <br> scenario | $\mathbf{M}_{\mathbf{w}}=\mathbf{4 . 5}$ | $\mathbf{M}_{\mathbf{w}}=\mathbf{5 . 1}$ | $\mathbf{M}_{\mathbf{w}}=\mathbf{5 . 7}$ |  |
| :--- | :---: | :---: | :---: | :---: |
| Distance, $\mathrm{d}(\mathrm{km})$ | $10<\mathrm{d}<100$ | $10<\mathrm{d}<100$ | $\mathrm{~d}>100$ |  |
| ${\text { Magnitude, } \mathrm{M}_{\mathrm{w}}}^{\mathrm{V}_{\mathrm{s}, 30}(\mathrm{~m} / \mathrm{s})}$ | $4<\mathrm{M}_{\mathrm{w}}<5$ | $4.5<\mathrm{M}_{\mathrm{w}}<6.0$ | $5.2<\mathrm{M}_{\mathrm{w}}<6.2$ |  |
| Fault type | Reverse, reverse/oblique/normal/strike slip |  |  |  |
| Source: the authors |  |  |  |  |

Table 2.
Accelerograms selected from PEER NGA-West 2 Ground Motion Database

| Record Sequence number (RSN) | Event | Year | Station | Mag. |
| :---: | :---: | :---: | :---: | :---: |
| 23 | San Francisco | 1957 | Golden Gate Park | 5.28 |
| 98 | Hollister-03 | 1974 | Gilroy Array \#1 | 5.14 |
| 146 | Coyote Lake | 1979 | Gilroy Array \#1 | 5.74 |
| 643 | Whittier Narrows-01 | 1987 | LA - Wonderland Ave | 5.99 |
| 680 | Whittier Narrows-01 | 1987 | $\begin{gathered} \text { Pasadena - CIT } \\ \text { Kresge Lab } \end{gathered}$ | 5.99 |
| 703 | Whittier Narrows-01 | 1987 | Vasquez Rocks Park | 5.99 |
| 1649 | Sierra Madre | 1991 | Vasquez Rocks Park | 5.61 |
| 1709 | Northridge-06 | 1994 | LA - Griffith Park Observatory | 5.28 |
| 1715 | Northridge-06 | 1994 | LA - Wonderland Ave | 5.28 |
| 4083 | Parkfield-02, CA | 2004 | Parkfield - Turkey Flat \#1 | 6 |
| 4312 | Umbria-03, Italy | 1984 | Gubbio | 5.6 |
| 2805 | Chi-Chi, Taiwan-04 | 1999 | KAU003 | 6.2 |

Source: the authors

Table 3.
Accelerograms selected from PEER NGA-West 2 Ground Motion Database (part 2)

| Record Sequence <br> number (RSN) | Mechanism | Rrup(km) | Vs30(m/s) | Scale <br> Factor |
| :---: | :---: | :---: | :---: | :---: |
| 23 | Reverse | 11.0 | 874.7 | 0.90 |
| 98 | strike slip | 10.5 | 1428.1 | 1.24 |
| 146 | strike slip | 10.7 | 1428.1 | 0.60 |
| 643 | Reverse | 27.6 | 1222.5 | 1.49 |
|  | Oblique |  |  |  |
| 680 | Reverse | 18.1 | 969.1 | 0.55 |
|  | Oblique |  |  |  |
| 703 | Reverse | 50.4 | 996.4 | 1.36 |
| 1649 | Oblique |  |  |  |
| 1709 | Reverse | 39.8 | 996.4 | 1.17 |
| 1715 | Reverse | 21.7 | 1015.9 | 1.34 |
| 4083 | Reverse | 17.1 | 1222.5 | 1.91 |
| 4312 | strike slip | 5.3 | 907.0 | 0.29 |
| 2805 | Normal | 15.7 | 922.0 | 0.99 |
|  | strike slip | 116.2 | 913.8 | 3.70 |

Source: the authors

The search results are shown in Table 2, 3, a total of twelve records were found. The response spectra of the scaled ground motions are presented in Fig. 4 including the geometric mean of the results. The process of searching and scaling the accelerograms returned a series of results that on average are compatible with the target response spectrum.

From Fig. 4, it can be observed that the scaled records either exceed or are equal to the design spectra only over a certain period range. In addition, the second part of the
application of the methodology consists of modifying the time series and makes them compatible with the target response spectra at all spectral periods is the spectral matching. The use of spectrum compatible time series for the seismic design of structures has been discussed by [33]. On one hand, spectrum compatible time series represent more than one earthquake and may be considered unrealistic when compared to typical earthquake response spectra. On the other hand, spectrum compatible time series has the advantage to reduce the number of time series necessary for an Engineering analysis [34], and the variability in spectral amplitude is considerably shortened.

Various methodologies have been proposed to obtain spectrum compatible time series [56,35,37-39]. Three basic approaches for spectral matching have been adopted in the different methods [33]: frequency-domain method, frequency-domain method with Random Vibration Theory (RVT), and time-domain method. The third approach is generally more complicated, however, produces a good convergence [40,41]. [42] implemented in RSPMatch program the Lilahanand and Tseng algorithm, a method based on the spectral matching in the time-domain. The adjustment wavelets used, however, produced drifts in the displacement and velocity time series and baseline corrections were required. The method was revised by [43] using new wavelets, to eliminate the drift in the adjustment wavelets to produce time series that do not require baseline correction.

The methodology applied in the current work and implemented in the software SeismoMatch [22] used the wavelets algorithm proposed by [33]. It has the advantage of matching the pseudo-acceleration response spectra of the time series with the target spectra, it does not introduce additional energy into the ground motion, preserves the characteristics of the original ground motion and eliminates the drift in the velocity and displacement time series.

As an example of the matching process in terms of the target response spectra, Fourier spectra, and Arias intensity, two earthquakes from Table 2 were selected. Arias Intensity (energy content) was included in the comparison of the


Figure 4. 5\% response spectra for the scaled accelerograms and target spectra.
Source: the authors.
matched records, considering its importance in seismic hazard analysis [44]. Fig. 5 shows the acceleration timehistories, Fourier spectra and Arias Intensities of the recorded and matched time series of earthquakes RSN 23 and RSN 4803. The Fourier amplitude maintains the characteristics of the initial time series. In addition, a good agreement was observed between the normalized Arias intensities of the initial (measured) and the modified (spectrally matched) time series. The same procedure was employed for all the records shown in Tables 2,3 and the spectral matching results are presented in Fig. 6.

From this methodology, it can be stated that nowadays it is relatively straightforward to have access to extensive databases of real strong-motion records. And the criteria for selecting the records are available in the literature [32] to obtain results representative of the seismicity of the area of interest. The spectrum compatible accelerograms obtained presented a satisfactory agreement also in terms of the frequency content and energy when compared to the original records.

### 2.2 Method 2: synthetic accelerograms

The second alternative consisted of the generation of artificial accelerograms based on the adaptation of a random process to a target spectrum. One of the most complete methodologies for obtaining ground motions is the specific barrier model developed by $[45,46]$. The model has been implemented and calibrated with extended databases of response spectral amplitudes from earthquakes of intraplate areas, interplate regions and regions of tectonic extension [47]. The advantage of this model is the use of relatively few parameters to generate records: (1) tectonic regime, (2) magnitude, (3) distance, and (4) soil/rock type. The method has been implemented


Figure 5. Example of matched accelerograms: acceleration time-histories, Fourier amplitude and normalized Arias intensity.
Source: the authors


Figure 6. Response spectra of matched accelerograms.
Source: the authors.
in the software SeismoArtif [48], in which the generation of the synthetic accelerogram starts from a Gaussian white noise that is multiplied by the [49] envelope shape and then adapted to a target spectrum using the Fourier Transform Method [50].

Fig. 7 presents a comparison between the Peak Ground Accelerations (PGA) obtained from the synthetic accelerograms versus distance and the curve obtained from the Ground Motion Prediction Equation (GMPE) developed by [51] and modified by [52], for the stable continental region in Central and Eastern North America (curve for a scaling factor equal to 1.00 in Fig. 7). Considering that Brazil is in an intraplate setting of low attenuation, [5] adopted [51] and [52] GMPE for the probabilistic hazard assessment in the continental margin of Southeastern Brazil. Additional curves were generated using Scaling Factors (SF) to cover the range of possible seismic attenuation in Brazil ( $\mathrm{SF}=0.4,0.75,1.00,1.33$ ). From Fig. 7, it can be observed a good agreement in terms of the PGA values obtained from the synthetic accelerograms and the curves obtained from GMPE.


Figure 7. Peak Ground Accelerations with distance: comparison between GMPE and synthetic accelerograms for a $4.5 M_{w}$ earthquake.
Source: the authors.

Table 4.
Earthquake scenarios for the generation of synthetic accelerograms

| Earthquake Scenario | $\mathrm{M}_{\mathrm{w}}=4.5$ | $\mathrm{M}_{\mathrm{w}}=5.1$ | $\mathrm{M}_{\mathrm{w}}=5.7$ |
| :---: | :---: | :---: | :---: |
| Distance $(\mathrm{km})$ | $10,50,100$ | $10,50,100$ | $100,200,400$ |
| Soil category |  | Very hard rock |  |
| Regime |  | Intraplate |  |

Source: the authors


Figure 8. Example of a synthetic accelerogram using the specific barrier model ( $\mathrm{Mw}=4.5$ ).
Source: the authors


Figure 9. Example of a synthetic accelerograms using the specific barrier model ( $\mathrm{Mw}=5.5$ ).
Source: the authors.

A series of artificial accelerograms were obtained using the criteria of the disaggregation analysis for earthquake scenarios as presented in Tables 1 and 4 presents in detail the criteria adopted for the generation of the synthetic accelerograms. Figs. 8 and 9 show examples of synthetic acceleration time-histories for two earthquakes of magnitudes 4.5 and $5.5 M_{w}$, for a rupture distance of 10 km .


Figure 10. Response spectra of matched synthetic accelerograms. Source: the authors.

Once obtained the synthetic accelerograms, the following stage consisted of performing a spectral matching based on the target spectra. The results are presented in Fig. 10. The algorithm was able to converge for all the synthetic accelerograms with slight variations for periods larger than one second.

### 2.3 Method 3: synthetic accelerograms using shape functions

A third approach to obtain synthetic accelerograms consisted of modifying a starting random process using an envelope shape [53-55,49] and a Power Spectral Density Function (PDSF). The method was proposed by [56], and it starts with the generation of steady-state motions (sinusoidal waves) with phase angles in the interval $(0,2 \pi)$ and amplitudes given by the PSDF (eq. 1).

$$
\begin{equation*}
a(t)=q(t) \sum_{i=1}^{n} A_{i} \sin \left(\omega_{i} t+\Phi_{i}\right) \tag{1}
\end{equation*}
$$

Where $q(t)$ is an envelope shape (or intensity function) that allows to resemble a real accelerogram; $A_{i}, \omega_{i}$ and $\Phi_{i}$ are respectively the amplitude, frequency, and the phase angle if the $i_{t h}$ sinusoidal wave. The generated ground motion is adapted to the target response spectrum using the Fourier Transform Method.

For the current analysis, it was adjusted an exponential envelope shape [54] defined by three parameters: duration, constants that modify the shape of the envelope ( $\alpha$ and $\beta$ ) and a parameter (A) that is function of the $\alpha$ and $\beta$ coefficients. Fig. 11 shows a representation of the envelope shape and the parameters involved in the calculation.

Given the lack of earthquake records in the continental margin of Southeastern Brazil by Ocean Bottom Seismometers (OBS's), two earthquakes of particular interest were selected to obtain measured envelope shapes. The first earthquake selected occurred on May 19th, 2012, in the town of Montes Claros, Brazil, a region in which seismicity is not a recent occurrence [57]. The second earthquake occurred on April 23rd, 2008, 125 km South of the city of São Vicente, on the coast of São Paulo State. Both earthquakes were recorded by many stations of the Brazilian seismic network


Figure 11. Envelope shape definition.
Source: the authors

Table 5.
Earthquakes recorded in Southeastern Brazil.

| Earthquake |  |  |  | Epicenter |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Earthquake | Date | Magnitude <br> $(\mathbf{M w )}$ | Latitude | Longitude | Focal <br> depth <br> $(\mathbf{k m})$ |  |
| Montes | $19 / 12 / 2012$ | 3.2 | -16.69 | -43.87 | 1.4 |  |
| Claros | $23 / 04 / 2008$ | 5.6 | -25.7 | -45.41 | 17 |  |
| São Vicente | 23 |  |  |  |  |  |

Source: the authors

Table6.
Earthquakes and stations that recorded the events.

| Earthquake |  |  | Station |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Earthquake | Date | Magnitude <br> (Mw) | Station | Latitude | Longitude |
| Montes <br> Claros <br> São | $19 / 12 / 2012$ | 3.2 | MCI9- <br> BL | -16.7 | -43.89 |
| Vicente | $23 / 04 / 2008$ | 5.6 | ESAR- <br> BL | -23.02 | -44.44 |

Source: the authors
and well-detailed information about the mentioned seismic events was presented by [3] and [57]. Tables 5 and 6 show information about the selected events and the location of the stations that registered the accelerograms used for the analysis.

Figs. 12 and 13 present the acceleration time-histories and the associated envelope shapes of the recorded motions. It can be observed that low accelerations were recorded from the São Vicente earthquake, this is due to the distance between the station and the epicenter (around 313.8 km ). The Montes Claros earthquake, on the other hand, was recorded by a station located at 1.44 km of the epicenter, then, relatively higher accelerations can be observed.

Using the envelope shapes of the two recorded accelerograms, the coefficients $\alpha$ and $\beta$ were defined to match an approximate envelope shape for the generation of the synthetic accelerogram (Fig. 14). The synthetic accelerogram and the response spectra are shown in Fig. 15.


Figure 12. Recorded earthquakes and envelope shapes: São Vicente (2008) earthquake.
Source: the authors.


Figure 13. Recorded earthquakes and envelope shapes: Montes Claros (2012) earthquake.

Source: the authors.


Figure 14. Adjusted envelope shape.
Source: the authors.


Figure 15. Artificial accelerogram using the adjusted envelope shape. Source: the authors.

## 3. Conclusions

This study presented three approaches to acquire spectrumcompatible accelerograms aimed to be applied for site response analyses or seismic design of structures in an area located in the Southeast region Brazil, specifically in the offshore Campos Basin. The first method presented, based on real accelerograms enables the obtention of accelerograms that are spectrum and energy-compatible accelerograms when compared with the historical records obtained from the PEER-NGA database. The second approach, based on seismological characteristics of the zone of interest, produced a good agreement in terms of Peak Ground Accelerations with distance obtained from the GMPE adopted for a stable continental region in Central and Eastern North America and adapted for the continental margin of Southeastern Brazil. The third alternative presented used information from real records in Southeastern Brazil, and well documented in the literature $[3,57]$, to generate a synthetic accelerogram utilizing an adjusted shape function to capture the characteristics of real ground motions.

Further analyses can be performed in the future with more data from earthquakes that occurred in Brazil: to refine the parameters for the generation of synthetic accelerograms or to adjust real records compatible also with the energy and frequency content.

## Acknowledgements

The authors thank the Rio de Janeiro State Research Foundation (FAPERJ) for its support to this research.

## References

[1] Assumpção, M., Ferreira, J., Barros, L., Bezerra, H., França, G., Barbosa, J. and Dourado, J., Intraplate seismicity in Brazil. In: Talwani, P., Ed., Intraplate Earthquakes Cambridge University Press., Cambridge, U.K., 2014, pp. 50-71. DOI: 10.1017/CBO9781139628921.004.
[2] Assumpção, M., Seismicity and stresses in the Brazilian passive margin. Bull Seismol Soc Am., 88(1), pp. 160-169, 1998a.
[3] Assumpção, M., Dourado, J.C., Ribotta, L.C., Mohriak, W.U., Dias, F.L. and Barbosa, J.R., The São Vicente earthquake of 2008 April and seismicity in the continental shelf off SE Brazil: further evidence for flexural stresses. Geophys J. Int., 187(3), pp. 1076-1088, 2011. DOI: 10.1111/j.1365-246 X.2011.05198.x.
[4] Cobbold, P.R, Meisling, K.E. and Mount, Van.S., Reactivation of an obliquely rifted margin, Campos and Santos Basins, Southeastern Brazil. AAPG Bulletin, 85(11), pp. 1925-1944, 2001. DOI: 10.1306/8626D0B3-173B-11D7-8645000102C1865D.
[5] Borges, R.G., Assumpçao, M.S., Almeida, M.C.F., Almeida. and M.S.S., Seismicity and seismic hazard in the continental margin of southeastern Brazil., J. Seismol., 24, pp. 1205-1224, 2020. DOI: 10.1007/s10950-020-09941-4.
[6] Amante, C. and Eakins B.W., ETOPO1 1 Arc-minute global relief model: procedures, data sources and analysis. NOAA Technical Memorandum NESDIS NGDC-24. National Geophysical Data Center, NOAA, 2009. DOI: 10.7289/V5C8276M
[7] Fetter, M., The role of basement tectonic reactivation on the structural evolution of Campos Basin, offshore Brazil: evidence from 3D seismic analysis and section restoration. Marine and Petroleum Geology, 26, pp. 873886, 2009. DOI: 10.1016/j.marpetgeo.2008.06.005.
[8] Dias, J.L., Scarton, J.C., Guardado, L.R., Esteves, F.R. and Carminatti, M., Aspectos da evolução tectono-sedimentar e da ocorrência de hidrocarbonetos na Bacia de Campos. In: Raja-Gabaglia, G.P., Milani, E.J., Eds., Origem e Evolução de Bacias Sedimentares. Petrobras, Rio de Janeiro, 1990, pp. 333-360.
[9] Chang, H.K., Kowsmann, R.O., Figueiredo, A.M.F. and Bender, A.A., Tectonics and stratigraphy of the East Brazil Rift system: an overview.

Tectonophysics, 213, pp. 97-138, 1992. DOI: 10.1016/0040-1951(92)90253-3
[10] Lima-Neto F.F. and Beneduzi, C., Using leakoff tests and acoustic logging to estimate insitu stresses at deep waters -Campos Basin. Extended Abstracts Volume, in: American Association of Petroleum Geologists International Conference \& Exhibition, Rio de Janeiro, Brazil, 1998, pp. 224-225.
[11] Bianchi, M.B., Assumpção, M., Rocha, M.P., Carvalho, J.M., Azevedo, P.A., Fontes, S.L., Dias, F.L., Ferreira, J.M., Nascimento, A.F., Ferreira, M.V. and Costa, I.S.L., The Brazilian seismographic network (RSBR): improving seismic monitoring in Brazil. Seismol. Res. Lett., 89(2A), pp. 452-457, 2018. DOI: 10.1785 /0220170227
[12] Almeida, A.A.D., Assumpção, M., Bommer, J.J., Drouet, S., Riccomini, C. and Prates, C.L.M., Probabilistic seismic hazard analysis for a nuclear power plant site in southeast Brazil, J. Seismol., 23, pp. 1-23, 2019. DOI: 10.1007/s10950-018-9755-8.
[13] Bianchi, M., Assumpção, M., Detzel, H.A., Carvalho, J.M., Rocha, M.P., Drouet, S., Fontes, S., Ferreira, J., Nascimento, M.A. and Veloso, J.A.V., The Brazilian seismographic network: historical overview and current status. Bull. Int. Seismol., Cent. 49, (1/6), pp. 70-90, 2015. DOI: 10.5281/zenodo. 998851
[14] Yi, J., Lam, N., Tsang, H.-H. and Au, F.T., Selection of earthquake ground motion accelerograms for structural design in Hong Kong. Advances in Structural Engineering, 23(10), pp. 2044-2056, 2020. DOI: 10.1177/1369433220906926
[15] Ahmadi G., Generation of artificial time-histories compatible with given response spectra: a review. SM Arch., 4(3), pp. 207-239, 1979.
[16] Lam, N., Wilson, J. and Hutchinson, G., Generation of synthetic earthquake accelerograms using seismological modelling: a review. Journal of Earthquake Engineering, 4(3), pp. 321-354. DOI: 10.1080/13632460009350374
[17] Iervolino I. and Cornell, A.C., Record selection for nonlinear seismic analysis of structures. Earthq Spectra, 21(3), pp. 685-713, 2005. DOI: 10.1193/1.1990199
[18] Ferreira, F., Moutinho, C., Cunha, Á. and Caetano, E., An artificial accelerogram generator code written in Matlab. Engineering Reports, 2(3), art. e12129, 2020. DOI: 10.1002/eng2.12129
[19] Barenberg, M.E., Inelastic response of a spectrum-compatible artificial accelerogram. Earthq Spectra., 5(3), pp.477-493, 1989. DOI: 10.1193/1.1585536
[20] Giaralis, A. and Spanos, P.D., Wavelets based response spectrum compatible synthesis of accelerograms - Eurocode application (EC8). Soil Dyn. Earthq. Eng., 29, pp. 219-235, 2009. DOI: 10.1016/j.soildyn.2007.12.002
[21] Hancock, J., Watson-Lamprey, J. and Abrahamson, N.A., An improved method of matching response spectra of recorded earthquake ground motion using wavelets. J. Earthq. Eng., 10, pp. 67-89, 2006. DOI: 10.1080/13632460609350629
[22] Seismosoft. SeismoMatch - A computer program for generation of artificial accelerograms, [online]. 2020. Available at: https://seismosoft.com/
[23] Mukherjee, S. and Gupta, V., Wavelet-based characterization of design ground motions. Earthquake Engineering \& Structural Dynamics. 31, pp. 1173-1190, 2002. DOI: 10.1002/eqe. 155.
[24] Vanmarcke, E.H. and Gasparini, D.A., Simulated earthquake groundmotions. Proceedings of the $4^{\text {th }}$ International Conference on SMIRT, K1/9, San Francisco, CA, USA, 1977.
[25] Taylor, C.A., EQSIM. A program for generating spectrum compatible earthquake ground acceleration time histories. Reference Manual. Bristol Earthquake Engineering Data Acquisition and Processing System, UK, 1989.
[26] Cacciola, P., A stochastic approach for generating spectrum compatible fully nonstationary earthquakes. Comput Struct, 88, pp. 889-901, 2010. DOI: 10.1016/j.compstruc.2010.04.009.
[27] Vanmarcke, E.H., Representation of earthquake ground motion: scaled accelerograms and equivalent response spectra, State-of-the-art for assessing earthquake hazards in the United States, Report 14, Miscellaneous Paper, S-73-1. US Army Corps of Engineers, Vicksburg, Mississippi, USA, 1979.
[28] Kramer, S.L., Geotechnical earthquake engineering. Prentice-Hall, New Jersey, USA, 1996.
[29] Bazzurro, P. and Cornell, C.A., Disaggregation of seismic hazard. Bull Seismol. Soc. Am., 89(2), pp. 501-520, 1999.
[30] Abrahamson, N.A., State of the practice of seismic hazard evaluation. In: Paper presented at the International Society for Rock Mechanics and Rock Engineering Symposium, Melbourne, Australia, November 19-24, 2000.
[31] Stewart, J.P., Chiou, S.-J., Bray, J.D., Graves, R.W., Somerville, P.G. and Abrahamson, N.A., Ground motion evaluation procedures for performancebased design, PEER Report 2001/09, Pacifc Earthquake Engineering Research Center, University of California, Berkeley, USA, 2001.
[32] Bommer, J.J. and Acevedo, A.B., The use of real earthquake accelerograms as input to dynamic analysis. Journal of Earthquake Engineering, 8(1), pp. 43-91, 2004. DOI: 10.1080/13632460409350521
[33] Al-Atik, L. and Abrahamson, N., An improved method for nonstationary spectral matching. Earthquake Spectra, 26(3), pp.601-617, 2010. DOI: 10.1193/1.3459159
[34] Bazzurro, P. and Luco, N., Do scaled and spectrum-matched near-source records produce biased nonlinear structural responses? In: Proc. of the $8^{\text {th }}$ U.S. National Conf. on Earthquake Engin., San Francisco, California, USA, 2006.
[35] Kaul, M.K., Spectrum-consistent time-history generation, ASCE J. Eng. Mech. EM4, pp.781-788, 1978. DOI: 0.1061/JMCEA3.0002379
[36] Gasparini, D.A. and Vanmarcke, E H., Simulated earthquake motions compatible with prescribed response spectra, Evaluation of Seismic Safety of Buildings Report No. 2, Department of Civil Engineering, MIT, Cambridge, Massachusetts, USA, 1979.
[37] Iyengar, R.N, and Rao, P., Generation of spectrum compatible accelerograms. Earthq. Eng. Struct. Dyn., 7(3), pp. 253-263, 1979. DOI: 10.1002/eqe. 4290070305
[38] Cacciola, P., Colajanni, P. and Muscolino, G., Combination of modal responses consistent with seismic input representation. J. Struct. Eng., 130(1), 2004. DOI: 10.1061/(ASCE)0733-9445(2004)130:1(47)
[39] Zentner, I. and Poirion, F., Enrichment of seismic ground motion databases using Karhunen-Loe`ve expansion. Earthq. Eng. Struct. Dyn., 41(14), pp. 1945-1957, 2012. DOI: 10.1002/eqe. 2166
[40] Lilhanand, K. and Tseng, W.S., Generation of synthetic time histories compatible with multiple-damping response spectra, SMiRT-9, Lausanne, K2/10, 1987.
[41] Lilhanand, K. and Tseng, W.S., Development and application of realistic earthquake time histories compatible with multiple damping response spectra, in: Ninth World Conf. Earth. Engin., Tokyo, Japan, (2), 1988, pp. 819-824.
[42] Abrahamson, N.A., Non-stationary spectral matching, Seismol. Res. Lett., 63, art. 30, 1992.
[43] Hancock, J., Watson-Lamprey, J., Abrahamson, N.A., Bommer, J.J., Markatis, A., McCoy, E. and Mendis, R., An improved method of matching response spectra of recorded earthquake ground motion using wavelets, J. Earthquake Eng., 10, pp. 67-89, 2006. DOI: 10.1080/13632460609350629
[44] Stafford, P.J., Berrill, J.B. and Pettinga, J.R., New predictive equations for Arias intensity from crustal earthquakes in New Zealand. J Seismol., 13(1), pp. 31-52, 2009a. DOI: 10.1007/s10950-008-9114-2
[45] Papageorgiou, A.S. and Aki, K., A Specific barrier model for the quantitative description of inhomogeneous faulting and the prediction of strong ground motion. Description of the model, Bull. Seism. Soc. Am., 73, pp. 693-722, 1983a.
[46] Papageorgiou, A.S. and Aki, K., Scaling law of far-field spectra based on observed parameters of the specific barrier model. PAGEOPH 123, pp. 353374 ,1985. DOI: 10.1007/BF00880736
[47] Hallodorsson, B. and Papageorgiou, A.S., Calibration of the specific barrier model to earthquakes of different tectonic regions Bull. Seismol. Soc. Am., 95(4), pp. 1276-1300, 2005. DOI: 10.1785/0120040157.
[48] Seismosoft. SeismoArtif - A computer program for generation of artificial accelerograms, [online]. 2020. Available at: https://seismosoft.com/
[49] Saragoni, G.R. and Hart, G.C., Simulation of artificial earthquakes. Earthq. Eng. Struct. Dyn., 2(3), pp.249-267, 1974. DOI: 10.1002/eqe. 4290020305.
[50] Mucciarelli, M., Masi A., Gallipoli, M.R., Harabaglia, P., Vona, M., Ponzo, F. and Dolce, M., Analysis of RC building dynamic response and soilbuilding resonance based on data recorded during a damaging earthquake. Bull. Seismol. Soc. Am., 94(5), pp.1943-1953, 2004. DOI: 10.1785/012003186.
[51] Toro, G.R., Abrahamson, N.A. and Schneider, J.F., Model of strong ground motions from earthquakes in central and eastern North America: best estimates and uncertainties Seismological Research Letters, 68(1), pp. 41-57, 1997. DOI: 10.1785/gssrl.68.1.41.
[52] Toro, G.R., Modification of the Toro et al. (1997) attenuation equations for large magnitudes and short distances. Risk Engineering, Inc.,2002.
[53] Hou, S., Earthquake simulation models and their applications. Tech. Rep., Department of civil engineering, Massachusetts Institute of Technology, USA, 1968.
[54] Liu, S.C., Autocorrelation, and power spectral density functions of the parkfield earthquake of june 27, 1966. Bull Seismol. Soc. Am., 59(4), pp.1475-1493, 1969.
[55] Jennings, P.C., Housner, G.W. and Tsai, N.C., Simulated earthquake motions. Tech. Rep., EERL California Institute of Technology, USA, 1968.
[56] Gasparini, D.A. and Vanmarcke, E.H., Simulated earthquake motions compatible with prescribed response spectra. Research Report R76-4. Department of Civil Engineering, Massachusetts Institute of Technology, Cambridge, MA, USA, 1976.
[57] Agurto-Detzel, H., Assumpção, M., Ciardelli, C., Farrapo, D., Barros, L.V. and França, G., The 2012-2013 Montes Claros earthquake series in the São Francisco Craton, Brazil: new evidence for non-uniform intraplate stresses in mid-plate South America, Geophysical Journal International, 200(1), pp. 216-226, 2015. DOI: 10.1093/gji/ggu333.
[58] Ancheta, T.D., Darragh, R.B., Stewart, J.P., Seyhan, E., Silva, W.J., Chiou, B.S.-J., Wooddell, K.E., Graves, R.W., Kottke, A.R., Boore, D.M., Kishida, T. and Donahue, J.L., NGA-West2 Database. Earthquake Spectra, 30(3), pp. 989-1005, 2014. DOI: 10.1193/070913EQS197M
C. Soriano, received the BSc. Eng. in Civil Engineering in 2011 from the Universidad Nacional de Colombia, Bogotá, Colombia. He also holds a MSc. in Civil Engineering from the Graduate School of Engineering of the Federal University of Rio de Janeiro, Brazil. Currently, he is a DSc. student in Civil Engineering with emphasis in Geotechnical Engineering at the Graduate School of Engineering of the Federal University of Rio de Janeiro, Brazil.
ORCID: 0000-0001-9530-0185
S. Tarazona, graduated in 2007 from Universidad Católica de Santa María in Peru. He holds a MSc. in Civil Engineering from the Pontifical Catholic University of Rio de Janeiro Brazil, and a DSc. from Graduate School of Engineering of the Federal University of Rio de Janeiro, Brazil. Currently, he is a researcher at COPPE/UFRJ Multidisciplinary Centrifuge Modeling Laboratory $\left(\mathrm{LM}^{2} \mathrm{C}\right)$ with interests related to the identification of the static and dynamic liquefaction triggers of tailings dams. ORCID: 0000-0001-5268-6487
M.C.F. Almeida, has a BSc. Eng. in Civil Engineering and Structures from the Federal University of Rio de Janeiro, Brazil, in 1975, a MSc. in Civil Engineering Structures from the Polytechnic of Central London, U.K., in 1984, and a PhD. in Civil Engineering Structures from UFRJ, Brazil, in 1997. She is currently an adjunct professor 3 from UFRJ, Brazil. She has experience in the field of civil engineering, with emphasis on static and dynamic structural analysis, acting mainly on the following themes: structural analysis, cracking, reinforced concrete, soil-structure interaction, seismic analysis, seismic risk analysis with Brazilian data, and technical standards.
ORCID: 0000-0002-3133-6098
M.S.S. Almeida, is currently a full professor in geotechnical engineering at COPPEUFRJ in Brazil. He did his undergraduate study in 1974 at the Civil Engineering Department of the School of Engineering, Federal University of Rio de Janeiro; and his MSc. from the same university in 1977. Prof. Almeida obtained his PhD. from Cambridge University, U.K., on stage constructed embankment in soft clays. He was a visiting researcher in many universities: Oxford University (1986-1989); Cambridge University (1993-1996); University of Western Australia (2002); LCPC, France (2011), and ETH, Switzerland (2012). He was also a postdoctoral research fellow at ISMES, Italy and NGI, Oslo, Norway in 1991-1992.
ORCID: 0000-0003-2230-397X
R.G. Borges, is BSc. in 2001 from the Lutheran University of Brazil. He holds MSc and DSc degrees in Civil Engineering from the Federal University of Rio de Janeiro, Brazil. He is currently a researcher at the Petrobras Research and Development Center (CENPES). He has experience in the fields of seismic hazards, soil-structure interaction, submarine slope stability physical modelling in geotechnical centrifuge, and salt rock Geomechanics.
ORCID: 0000-0001-7623-0337


[^0]:    How to cite: Soriano. C., Tarazona, S.F., Almeida, M.C.F., Almeida, M.S.S. and Borges, R.G., Selection of earthquake ground motion accelerograms for the continental margin

